

## REVEL: A model for Recent plate velocities from space geodesy

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[1] We present a new global model for Recent plate velocities, REVEL, describing the relative velocities of 19 plates and continental blocks. The model is derived from publicly available space geodetic (primarily GPS) data for the period 1993–2000. We include an independent and rigorous estimate for GPS velocity uncertainties to assess plate rigidity and propagate these uncertainties to the velocity estimates. The velocity fields for North America, Eurasia, and Antarctica clearly show the effects of glacial isostatic adjustment, and Australia appears to depart from rigid plate behavior in a manner consistent with the mapped intraplate stress field. Two thirds of tested plate pairs agree with the NUVEL-1A geologic (3 Myr average) velocities within uncertainties. Three plate pairs (Caribbean–North America, Caribbean–South America, and North America–Pacific) exhibit significant differences between the geodetic and geologic model that may reflect systematic errors in NUVEL-1A due to the use of seafloor magnetic rate data that do not reflect the full plate rate because of tectonic complexities. Most other differences probably reflect real velocity changes over the last few million years. Several plate pairs (Arabia–Eurasia, Arabia–Nubia, Eurasia–India) move more slowly than the 3 Myr NUVEL-1A average, perhaps reflecting long-term deceleration associated with continental collision. Several other plate pairs, including Nazca–Pacific, Nazca–South America and Nubia–South America, are experiencing slowing that began ~25 Ma, the beginning of the current phase of Andean crustal shortening. *INDEX TERMS*: 1243 Geodesy and Gravity: Space geodetic surveys; 8107 Tectonophysics: Continental neotectonics; 8150 Evolution of the Earth: Plate boundary—general (3040); 8158 Evolution of the Earth: Plate motions—present and recent (3040); *KEYWORDS*: Plate tectonics, geodesy, GPS, global plate model, present-day, REVEL

### 1. Introduction

[2] The present-day velocities of the Earth's lithospheric plates are an important kinematic boundary condition for many geologic and geophysical studies, including regional neotectonics, seismogenic zone processes, and earthquake hazards. Currently, the most comprehensive picture of geologically young plate motion comes from the global geologic model NUVEL-1A [DeMets *et al.*, 1990, 1994], a significant update of earlier global models [Chase, 1972, 1978; Minster *et al.*, 1974; Minster and Jordan, 1978]. NUVEL-1A is based in large part on mid-ocean ridge spreading rates dated from magnetic anomaly 2A (~3 Ma, or mid-Pliocene time) and thus describes relative plate velocities averaged over Pliocene to Recent time. NUVEL-1A is a robust model based on a large data set, but it nevertheless has some deficiencies. First, some smaller plates are necessarily omitted because of lack of data. Second, the geologic model may be biased due to poor or insufficient kinematic data, e.g., North America–Pacific relative motion [DeMets, 1995; DeMets and Dixon, 1999] and motion of the Caribbean plate relative to North and South America [Dixon *et al.*, 1998; DeMets *et al.*, 2000; Weber *et al.*, 2001; Perez *et al.*, 2001]. Third, the 3-Myr average velocity predicted by NUVEL-1A or any other geologic model may yield biased estimates of

present-day velocity for some plate pairs because the plates are speeding up, slowing down, or changing direction, e.g., Nazca–South America [Norabuena *et al.*, 1998, 1999; Angermann *et al.*, 1999].

[3] Space geodesy has the potential to measure relative plate velocities directly over periods of just a few years, as demonstrated by satellite laser ranging (SLR) [Smith *et al.*, 1990; Robbins *et al.*, 1993; Cazenave *et al.*, 1993; Sengoku, 1998], very long baseline interferometry (VLBI) [Argus and Gordon, 1990; Robaudo and Harrison, 1993; Ryan *et al.*, 1993; Sato, 1993]; Doppler Orbitography and Radiopositioning Integrated by Satellite (DORIS) [Cazenave *et al.*, 1992; Soudarin and Cazenave, 1993, 1995; Cretaux *et al.*, 1998] and the Global Positioning System (GPS) [Dixon *et al.*, 1991a; Dixon, 1993; Argus and Heflin, 1995; Larson *et al.*, 1997; Dixon and Mao, 1997]. Most studies to date suggest that the great majority of plate velocities estimated from space geodesy are consistent with the NUVEL-1A model within 95% confidence. However, uncertainties in the geodetic estimates have been large enough that important differences may have been missed. The true uncertainty of space geodetic data has also been difficult to quantify. Therefore it has been difficult to address an important tectonophysical problem, namely, the extent to which individual plate velocities may be changing over the last few million years and to what extent such changes, if they occur, can be understood in terms of simple plate-driving forces.

[4] New space geodetic data and new analytical techniques now permit a significant refinement of our description of present-day plate motion. In this paper we present a comprehensive velocity model for most major and several minor plates and

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continental blocks, based primarily on GPS. Our study differs from previous global studies in several respects:

1. A very large geodetic data set is now publicly available through the efforts of many individuals, institutions, and geodetic agencies, permitting a more accurate and more comprehensive geodetic plate motion model. This data set represents both a large number of sites, giving generally good geographic distribution, and long, nearly continuous time series at many individual sites, resulting in precise site velocity estimates. Most major plates now have at least two GPS sites, the minimum number to determine a plate's angular velocity with space geodesy.

2. We include velocity estimates for the Amuria, Anatolia, Caribbean, Nubia, Okhotsk, Philippine, Sierra Nevada, Somalia, South China, and Sunda plates or continental blocks. While these and other plates and continental blocks have been the subject of earlier local studies, most previous global geologic and geodetic plate motion models have omitted one or more of these plates because of sparse data or have approximated Nubia (west Africa) by combining it with Somalia (east Africa).

3. We incorporate a rigorous, independent estimate for GPS velocity errors. This permits simple, objective tests of whether the GPS site velocity for a given location is consistent with rigid plate behavior and whether plate motions averaged over the last few years differ significantly from motions averaged over the last few million years.

[5] The velocity predictions of geologic plate motion models are sometimes termed "present-day" or "current" because they describe geologically young plate motion, derived from the youngest easily identified magnetic anomaly, typically 2A (~3 Ma, or mid-Pliocene) [Minster and Jordan, 1978; DeMets et al., 1990, 1994] to, in some cases, anomaly 1N (~0.8 Ma, or mid-Pleistocene) [DeMets, 1995; Conder and Forsyth, 2000]. Our geodetic plate motion model is derived from data over a very different time span, roughly the last decade. It is probably representative of plate motions over the Holocene or Recent epoch (last ~10,000 years) and possibly the late Pleistocene epoch (last few hundred thousand years) provided that we account for, or avoid, short-term strain effects related to the seismic cycle and isostatic effects associated with the last glacial cycle. The former can impact site velocities in the vicinity of active plate boundary zone faults, while the latter may impact site velocities in parts of North America, Eurasia, Greenland, and Antarctica. To emphasize the time span over which we believe our model to be valid, we have termed it "REVEL" (for Recent velocities) with the suffix 2000, to indicate the last year of data included in the model. We expect that the model can be improved significantly in subsequent years as additional data are added and time series lengthen.

## 2. Data Analysis

[6] Uncertainty in the positions of the GPS satellites is a major error source for the coordinate time series used in this study. By 1993 the global tracking network for GPS satellites became sufficiently robust to produce more accurate satellite ephemerides (satellite positions as a function of time) compared to earlier periods. Our data span the time period 1 January 1993 to 31 December 2000. To ensure consistency, all data were reanalyzed specifically for this study, resulting in a uniform set of site velocity and error estimates. The great majority of sites used in this study are continuous, in the sense that they record data for 20–24 hours per day, typically for at least 300 days per year (Table 1). For more than 98% of the continuous sites presented here, all known existing data acquired after 1 January 1993 were processed (for the remaining sites, there may be additional data that were not available to us). In some cases (the Caribbean and Philippine plates and the Sierra Nevada block), continuous site distribution is limited, and we have augmented these data with data from

"episodic" sites occupied periodically, usually every year or two. A total of ~345,000 station days of data were analyzed for this study, most of which lie in stable plate interiors and are used to estimate plate velocities (Table 1). The data were analyzed at the University of Miami, generally following procedures outlined by Dixon et al. [1993, 1997], although some of these procedures have been updated considerably. Salient points are listed here:

1. We use GIPSY/OASIS II, Release 5.0 software developed at the Jet Propulsion Laboratory (JPL) and nonfiducial satellite orbit and clock files provided by JPL [Zumberge et al., 1997].

2. We use the ionosphere-free combination of both undifferenced carrier phase and P code pseudorange data (data weights of 1 cm and 1 m, respectively), typically recorded at 30-s intervals, with a 5-min decimation rate.

3. We created a comprehensive antenna height and type change file going back to 1993 for all sites, and used the relative dome-free models of Mader [1999] for first-order antenna phase center corrections.

4. To improve the definition of horizontal atmospheric gradients and to reduce their effect on the final position estimates, we use an elevation angle cutoff of 10° (where data are available) and the horizontal gradient model of Bar-Sever et al. [1998]. We use the same random walk model for the zenith delay as Bar-Sever et al. (3 mm/√h), but a looser constraint for the two orthogonal horizontal gradients (5 mm/√h). We use the mapping function of Niell [1996], which describes how the average atmospheric path delay varies as a function of elevation angle.

5. All sites are corrected for ocean tidal loading, using the NLOADF program of Agnew [1997], and the Schwiderski [1980] ocean tide model.

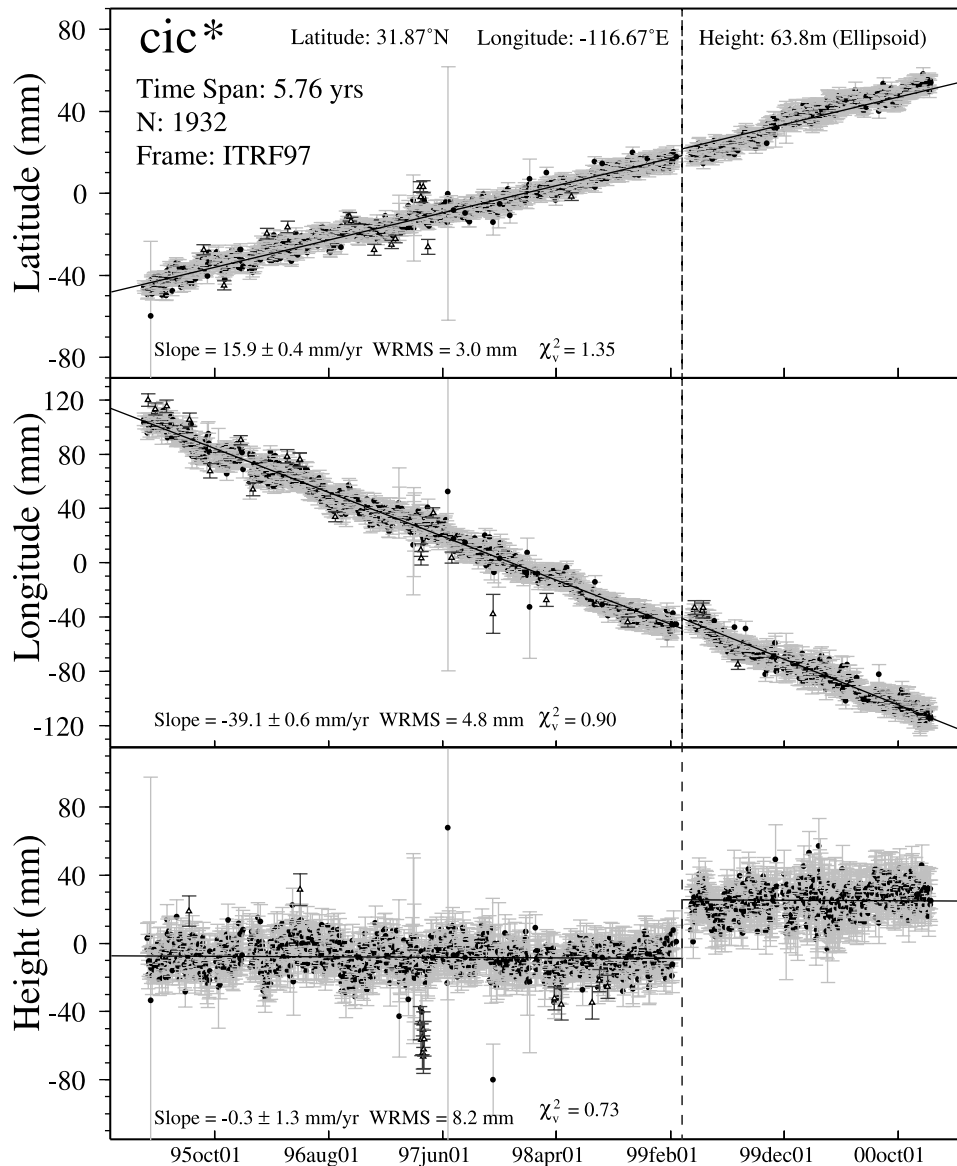
6. Carrier phase cycle ambiguities are estimated, not fixed.

7. We estimated offset parameters at the date of each change of antenna height or model to correct for second-order effects such as inaccuracy of phase center estimates and incorrect records of antenna height changes; these can exceed 1 cm in the vertical component, but generally are a few millimeters or less for horizontal components. Offsets were also calculated for antenna dome changes.

8. Daily position estimates are generated with loose constraints [Heflin et al., 1992; Blewitt et al., 1992], then transformed to International Terrestrial Reference Frame (ITRF)-97 [Boucher et al., 1999] using up to 51 collocated sites whose positions are defined in ITRF-97. The number of collocated sites depends on station availability. For 1998 and later, this usually exceeds 45, but in 1993, for example, it was typically 15–20. Transformation parameters are computed each day; hence each daily position estimate is essentially independent.

9. Position estimates with formal errors >9.9 m are removed from the database (typically <2 days per site per year).

[7] Velocity estimates are based on a weighted least squares line fit to the daily position estimates, including the offset parameters described above (Figure 1). Outliers, defined as points that lie off the best fit line by more than 3 times the formal error, are flagged, but not removed from the database, and are not used in the line fit (the fitting and outlier definition are done iteratively). Sites listed in upper case letters in Table 1 and in figures (e.g., STJO) lie on the stable plate and are used to define the rigid plate angular velocity. Sites in lowercase (e.g., cic\*, algo) are not used in the rigid plate definition because they may lie in the deforming boundary zone or may be affected by glacial isostatic adjustment or other nonrigid process. In some cases a station was moved or offset during the 1993–2000 period of this study. Where possible, we used publicly available vector tie information ("site ties") to link the two time series. If vector tie information was unavailable, we used a procedure similar to that described in point 7 above to tie the two time series together, in effect estimating the vector tie along with the slope (velocity) parameter of interest (Figure 1 shows an example). In such cases we are estimating three parameters (slope, intercept, and offset) rather



**Figure 1.** Example coordinate time series for site *cic\** (Table 1), a composite time series composed of *Cice* prior to 1 March 1999 and *Cic1* thereafter. Note offset estimate at date of site change (vertical dashed line) after site tie correction. Site velocity is given by slope of weighted least squares line fit (thin solid line) through the data (solid circles), excluding outliers (open triangles). Site velocity error is the uncertainty of the slope estimate, accounting for uncorrelated and time-correlated noise, total time span of observations, and total number of observations.

than the normal two parameters for a straight line. This allows us to take advantage of the reduction of velocity uncertainty associated with longer time series. Site names for such linked time series are indicated by asterisks in Table 1 and in the figures; e.g., *cic\** represents the combination of time series from sites *Cice* and *Cic1* (Figure 1). Velocities for the original sites are also listed in Table 1 immediately preceding the combined solution with only the first letter capitalized. In a few cases, time series have obvious offsets, presumably associated with an equipment or other significant change, even though no change is reported. In cases where such offsets are large enough to affect the velocity significantly, we have estimated the offset following the procedure outlined above. These sites are noted in Table 1.

[8] Our velocity error estimates account for both white (uncorrelated) and colored (time-correlated) noise [Langbein and Johnson, 1997] and are modeled following Mao *et al.* [1999]. This approach is based on a numerical analysis of 23 globally

distributed GPS sites with time series spanning the period 1994–1997. The error model avoids any assumptions concerning the fit of GPS data to a particular geophysical model (e.g., rigid plate model) and thus provides independent error estimates. To account for improvements in analytical techniques since 1997 and for site-specific effects, we exploit the correlation between WRMS (the weighted root mean square scatter of the daily position estimates about a best fit straight line) and white and flicker noise amplitudes observed in the data of Mao *et al.* [1999], as outlined by Dixon *et al.* [2000a]. Random walk noise is assumed to be zero. When this error model is applied to our data set of 64 site velocities from the stable interior of North America, we obtain a  $\chi^2$  per degree of freedom of 1.05 for the rigid plate model, close to the expected value of 1.0. This suggests that the error model is reasonable and that the region sampled by these data can be assumed rigid within data uncertainty. This is discussed in more detail in section 4.1.

**Table 1.** Site Positions and Velocities Relative to ITRF-97

Site <sup>a</sup>	Position <sup>b</sup>		$\Delta T$ , <sup>c</sup> years	Total Data <sup>d</sup>	Velocity <sup>e</sup>			WRMS <sup>f</sup>			Rate Res. <sup>g</sup>	
	°N	°E			North	East	Vertical	N	E	V	N	E
Amuria (Am)												
Taej <sup>h</sup>	36.37	127.37	3.32	1185	-17.8 ± 0.9	27.9 ± 1.2	3.8 ± 2.7	3.7	5.9	11.3	-	-
Daej <sup>h</sup>	36.40	127.37	1.79	619	-12.4 ± 1.6	25.6 ± 2.1	-1.0 ± 4.8	3.5	5.3	10.2	-	-
DAE* <sup>i</sup>	36.40	127.37	5.12	1804	-17.0 ± 0.6	27.6 ± 0.8	3.0 ± 1.8	4.0	5.8	11.1	-1.1	-0.2
SUWN	37.28	127.05	3.07	920	-14.2 ± 0.9	27.8 ± 1.2	1.9 ± 2.7	3.2	5.2	9.2	1.7	0.2
VLAD	43.20	131.93	4.47	1363	-15.7 ± 0.8	24.8 ± 0.8	-1.5 ± 2.0	4.4	5.4	11.0	0.3	0.1
bjfs	39.61	115.89	1.20	419	-13.8 ± 2.4	29.9 ± 2.8	-3.9 ± 7.5	3.4	4.9	11.0	1.3	1.1
irkt	52.22	104.32	5.27	1867	-9.1 ± 0.7	26.5 ± 0.7	-1.8 ± 1.7	4.3	5.4	10.6	4.6	-0.3
xian	34.37	109.22	3.48	991	-14.7 ± 0.9	34.5 ± 1.1	0.1 ± 2.5	3.8	5.2	10.0	-0.4	3.4
Anatolia (At)												
Ankr <sup>h</sup>	39.89	32.76	5.52	1580	10.3 ± 0.7	0.9 ± 1.0	-0.9 ± 1.5	4.9	7.5	9.4	0.0	0.0
ANK* <sup>j</sup>	39.89	32.76	4.14	1280	8.1 ± 0.7	4.8 ± 0.7	-3.7 ± 1.9	3.4	4.4	8.6	-0.2	0.2
7580 <sup>k</sup>	37.38	33.19	10	-	12.0 ± 4.0	5.2 ± 3.9	-5.2 ± 4.0	-	-	-	3.1	-3.7
7585 <sup>k</sup>	39.80	34.81	10	-	17.7 ± 4.3	-3.2 ± 4.3	-3.3 ± 4.3	-	-	-	6.8	-8.2
7589 <sup>k</sup>	39.89	32.76	10	-	8.0 ± 2.0	5.3 ± 1.7	21.1 ± 2.0	-	-	-	-0.3	0.8
Antarctica (An)												
CASI	-66.28	110.52	6.49	1839	-11.6 ± 0.7	3.7 ± 0.6	6.4 ± 1.5	5.2	5.5	12.4	0.2	0.2
DAV1	-68.58	77.97	6.49	1849	-6.6 ± 0.7	-2.1 ± 0.7	4.0 ± 1.5	5.2	6.3	12.2	0.3	0.4
KERG	-49.35	70.26	6.13	1971	-5.3 ± 0.7	5.2 ± 0.8	6.5 ± 1.5	5.1	7.1	11.7	0.1	0.3
MAW1	-67.60	62.87	7.00	1830	-3.6 ± 0.7	-3.5 ± 0.7	3.9 ± 1.3	5.8	7.2	10.2	0.2	-0.1
MCM4	-77.84	166.67	5.93	2060	-12.2 ± 0.6	10.6 ± 0.7	4.6 ± 1.9	4.5	5.7	15.5	-0.9	-0.5
SYOG	-69.01	39.58	4.00	1164	0.3 ± 1.0	-4.4 ± 1.2	5.8 ± 2.2	5.0	6.8	10.4	-1.2	0.1
VESL	-71.67	357.16	2.39	648	8.2 ± 1.8	-4.7 ± 1.4	1.1 ± 3.6	5.0	4.8	9.8	-1.7	-3.3
ohig	-63.32	302.10	5.70	1330	10.1 ± 0.9	14.7 ± 0.8	8.7 ± 1.8	6.1	6.5	12.9	-2.6	2.3
palm	-64.78	295.95	2.46	851	14.1 ± 1.9	11.7 ± 1.5	3.3 ± 3.4	5.7	5.3	9.5	1.8	-1.5
Arabia (Ar)												
BAHR	26.21	50.61	4.52	1592	26.9 ± 0.6	29.5 ± 0.7	-0.9 ± 1.8	3.0	4.9	9.1	0.2	-0.6
KATZ	33.00	35.69	4.18	590	19.3 ± 0.7	22.7 ± 1.0	3.9 ± 2.4	3.2	5.3	11.4	-0.2	1.1
7832 <sup>k</sup>	24.91	46.40	5	-	20.3 ± 8.2	25.9 ± 8.5	-4.3 ± 8.4	-	-	-	-4.6	-4.3
Australia (Au)												
ALIC	-23.67	133.89	6.45	1574	57.0 ± 0.4	33.9 ± 0.7	2.0 ± 1.5	3.1	6.0	11.8	0.1	-0.4
CEDU	-31.87	133.81	6.63	915	57.4 ± 0.5	31.5 ± 0.6	3.5 ± 1.4	3.4	5.2	10.0	0.5	0.5
DARW	-12.84	131.13	6.61	972	57.4 ± 0.5	38.4 ± 1.0	2.2 ± 1.6	3.4	8.2	12.2	0.2	0.2
HOB2	-42.80	147.44	6.49	1873	54.2 ± 0.5	15.5 ± 0.6	2.5 ± 1.4	3.7	5.6	10.7	0.2	-1.1
JAB1	-12.66	132.89	3.40	501	58.2 ± 0.8	35.3 ± 1.4	-1.9 ± 2.8	3.1	6.1	10.5	1.2	-2.6
KARR	-20.98	117.10	6.44	1576	56.7 ± 0.4	43.3 ± 0.7	1.0 ± 1.4	3.2	6.5	10.9	0.6	2.1
PERT	-31.80	115.89	7.36	2387	55.5 ± 0.4	42.1 ± 0.6	0.4 ± 1.3	3.3	6.3	11.8	-0.4	1.8
TID2	-35.40	148.98	5.00	1386	53.4 ± 0.5	19.5 ± 0.7	3.6 ± 1.7	3.2	4.9	10.0	-0.1	-1.4
TIDB	-35.40	148.98	8.00	2468	53.3 ± 0.5	19.4 ± 0.6	5.6 ± 1.2	4.3	6.7	12.0	-0.2	-1.5
TOW2	-19.27	147.06	5.96	1500	53.5 ± 0.5	30.5 ± 0.8	3.5 ± 1.6	3.5	6.8	11.1	-0.7	-1.1
YAR1	-29.05	115.35	8.00	2718	54.9 ± 0.4	41.9 ± 0.6	3.8 ± 1.2	3.5	6.3	11.7	-0.9	0.8
auck	-36.60	174.83	5.29	1885	37.5 ± 0.5	5.1 ± 0.7	3.9 ± 1.5	3.4	5.7	8.9	-1.8	-2.3
coco <sup>l</sup>	-12.19	96.83	4.55	1363	47.7 ± 0.7	43.0 ± 1.3	0.3 ± 2.2	3.9	8.5	12.4	-1.1	-2.2
dgar	-7.27	72.37	4.60	1556	29.8 ± 0.7	46.0 ± 1.4	2.5 ± 2.2	4.1	8.8	13.5	-2.3	0.5
hyde	17.42	78.55	5.47	268	32.2 ± 0.8	42.1 ± 1.5	2.1 ± 2.5	4.4	9.5	15.2	-4.8	17.0
iisc	13.02	77.57	5.98	1611	31.8 ± 0.6	42.4 ± 1.0	-0.1 ± 1.7	3.9	8.2	12.8	-4.5	13.5
mald <sup>l</sup>	4.19	73.53	1.39	485	30.1 ± 2.8	41.5 ± 4.2	8.6 ± 6.5	4.7	8.2	11.0	-3.0	5.1
noum	-22.27	166.41	3.00	986	43.1 ± 0.9	21.8 ± 1.6	1.8 ± 2.9	3.1	6.7	10.2	-1.9	-1.8
Caribbean (Ca)												
AVES	15.67	-63.62	3.87	29	11.5 ± 0.9	13.2 ± 2.1	5.0 ± 3.4	2.5	7.0	9.4	1.1	0.6
BARB	13.09	300.39	3.09	569	12.6 ± 1.0	13.1 ± 1.8	-0.7 ± 3.3	3.3	7.3	11.8	0.2	-0.3
CRO1	17.76	295.42	6.96	1828	10.7 ± 0.4	10.5 ± 0.7	2.5 ± 1.5	3.5	6.8	13.1	0.1	-0.6
ROJO	17.90	-71.67	6.25	21	6.2 ± 0.9	11.5 ± 1.9	-5.5 ± 3.7	3.8	9.0	18.4	0.8	-1.7
SANA	12.52	-81.73	6.25	29	4.8 ± 0.9	14.1 ± 1.2	0.8 ± 3.8	4.1	6.2	20.8	1.1	1.0
TDAD	10.68	-61.40	4.05	7	10.7 ± 1.4	15.9 ± 3.2	5.9 ± 4.1	3.2	8.1	9.4	1.5	-1.1
isab	18.47	-67.05	4.18	26	9.3 ± 1.4	11.1 ± 2.0	-0.8 ± 7.8	4.2	7.0	30.8	0.4	-0.4
pur3	18.46	292.93	3.56	1096	9.5 ± 0.8	7.5 ± 1.4	0.2 ± 2.9	3.2	7.0	13.7	-0.2	-3.2
Eurasia (Eu)												
ARTU	56.43	58.56	1.40	388	4.8 ± 1.6	24.5 ± 1.6	-0.9 ± 5.6	2.6	3.3	8.2	-0.2	-0.8
BOGO	52.48	21.04	4.00	1293	12.7 ± 0.5	22.2 ± 0.5	0.9 ± 2.0	2.4	3.0	8.8	0.1	0.8
BOR1	52.28	17.07	6.26	2221	12.5 ± 0.3	20.4 ± 0.4	-1.9 ± 1.2	2.6	3.8	8.1	-0.6	-0.4
GLSV	50.36	30.50	2.85	948	11.6 ± 0.9	22.4 ± 0.8	0.0 ± 2.7	2.9	3.5	8.1	0.6	-1.0
GOPE	49.91	14.79	5.30	1826	13.0 ± 0.4	21.6 ± 0.5	-4.9 ± 1.7	2.6	4.0	10.9	-0.4	0.7
JOZE	52.10	21.03	7.36	2577	12.5 ± 0.3	21.3 ± 0.4	-5.2 ± 1.2	3.0	4.1	10.8	-0.1	-0.2
KSTU	55.99	92.79	2.92	713	-5.6 ± 1.1	25.3 ± 1.3	-0.4 ± 2.9	3.6	5.2	9.5	-1.7	-0.4
LAMA	53.89	20.67	6.03	1844	12.8 ± 0.4	20.9 ± 0.5	-2.6 ± 1.3	3.1	4.4	8.2	0.2	-0.1



**Table 1.** (continued)

Site <sup>a</sup>	Position <sup>b</sup>		$\Delta T$ , <sup>c</sup> years	Total Data <sup>d</sup>	Velocity <sup>e</sup>			WRMS <sup>f</sup>			Rate Res. <sup>g</sup>	
	°N	°E			North	East	Vertical	N	E	V	N	E
MDVO <sup>l</sup>	56.03	37.22	5.84	2030	10.5 ± 0.5	23.4 ± 0.5	-5.4 ± 1.7	3.7	4.3	12.8	0.7	0.3
NYAL	78.93	11.87	7.97	2405	14.5 ± 0.4	10.4 ± 0.4	4.8 ± 1.1	3.5	4.1	10.0	0.8	-0.3
POTS	52.38	13.07	6.25	2131	13.4 ± 0.3	20.0 ± 0.4	-0.1 ± 1.2	2.4	3.5	7.7	-0.2	0.0
TIXI	71.63	128.87	2.23	764	-11.8 ± 1.3	16.2 ± 1.2	2.7 ± 3.7	3.6	3.9	9.4	-0.1	-0.5
WROC	51.11	17.06	3.66	533	13.9 ± 0.6	20.6 ± 0.6	-0.7 ± 2.1	2.5	3.1	7.0	0.8	-0.4
WTZR	49.14	12.88	4.98	1770	13.5 ± 0.4	20.5 ± 0.5	-0.6 ± 1.4	2.4	3.4	7.2	-0.1	-0.3
ZWEN	55.70	36.76	5.74	1841	9.6 ± 0.5	23.8 ± 0.5	-4.5 ± 1.5	3.1	4.4	9.5	-0.3	0.7
bili	68.08	166.44	1.33	369	-21.1 ± 2.0	11.4 ± 1.6	1.2 ± 5.9	3.1	3.1	8.1	-6.1	1.9
brst	48.38	355.50	2.16	528	15.8 ± 0.9	13.9 ± 1.3	-4.2 ± 3.4	2.2	4.0	7.0	0.9	-3.8
brus	50.80	4.36	6.69	2365	13.7 ± 0.3	18.0 ± 0.4	0.0 ± 1.1	2.5	3.8	7.5	-0.7	-0.8
casc	38.69	350.58	3.67	980	15.2 ± 0.6	15.6 ± 0.9	0.6 ± 2.1	2.5	4.6	7.7	0.2	-3.9
geno	44.42	8.92	3.77	937	15.6 ± 0.6	19.8 ± 0.7	-4.3 ± 2.3	2.6	3.9	9.3	1.6	-1.4
hers	50.87	0.34	7.91	2417	15.0 ± 0.4	15.7 ± 0.5	0.1 ± 1.1	3.3	6.0	9.5	0.3	-2.2
hofn	64.27	344.80	3.24	1093	14.7 ± 0.7	11.4 ± 0.8	4.0 ± 2.3	2.7	3.8	7.7	-0.3	1.5
irkt	52.22	104.32	5.27	1867	-9.1 ± 0.7	26.5 ± 0.7	-1.8 ± 1.7	4.3	5.4	10.6	-2.4	0.9
kiru	67.86	20.97	7.48	2599	13.4 ± 0.5	15.9 ± 0.6	4.7 ± 1.3	4.2	5.9	12.5	0.8	-0.9
kit3	39.13	66.89	6.24	1631	2.3 ± 0.5	28.6 ± 0.7	-2.6 ± 1.4	3.8	6.2	10.1	-0.6	0.4
kosg	52.18	5.81	8.00	2804	14.7 ± 0.3	17.7 ± 0.3	-1.1 ± 0.9	2.5	4.0	7.5	0.4	-0.9
mans	48.02	0.16	2.74	841	15.6 ± 0.7	16.7 ± 1.0	-2.7 ± 2.5	2.2	4.0	6.2	0.9	-2.0
mets	60.22	24.40	8.00	2684	11.4 ± 0.3	19.8 ± 0.4	2.7 ± 1.0	3.3	4.9	8.5	-0.7	-0.1
mlvl <sup>l</sup>	48.84	2.59	2.79	801	13.8 ± 1.7	16.8 ± 1.1	-8.2 ± 2.8	5.5	4.5	8.1	-0.7	-2.2
noto	36.88	14.99	4.88	1657	16.4 ± 0.5	21.9 ± 0.6	0.2 ± 1.7	3.2	4.6	9.0	3.0	-1.7
onsa	57.40	11.93	8.00	2779	13.0 ± 0.3	15.9 ± 0.4	1.7 ± 0.9	2.7	4.2	7.5	-0.7	-2.4
penc	47.79	19.28	4.67	1525	12.6 ± 0.5	22.6 ± 0.5	-7.3 ± 2.1	2.6	3.7	12.5	-0.2	0.4
pol2	42.68	74.69	5.61	1632	2.1 ± 0.5	28.3 ± 0.7	0.2 ± 1.4	3.5	5.2	8.1	1.3	0.2
reyk	64.14	338.04	4.95	1676	18.8 ± 0.6	-11.4 ± 0.6	-3.4 ± 1.5	3.4	4.4	7.9	4.0	-19.8
sjdv	45.88	4.68	3.20	1045	14.8 ± 0.6	18.6 ± 0.7	-0.6 ± 2.3	2.2	3.5	7.5	0.4	-1.5
sofi	42.56	23.39	3.48	972	10.4 ± 0.7	25.3 ± 1.0	-2.7 ± 2.4	2.9	4.7	9.3	-1.8	1.4
sumk	44.21	74.00	3.22	1002	-1.6 ± 0.9	27.8 ± 0.9	4.1 ± 2.8	3.3	4.4	10.7	-2.6	-0.1
toul	43.56	1.48	3.84	1228	15.0 ± 0.5	18.5 ± 0.7	-2.3 ± 1.9	2.3	4.0	7.2	0.4	-1.6
trom	69.66	18.94	7.99	2139	15.5 ± 0.4	14.0 ± 0.4	1.7 ± 1.0	4.1	5.0	8.8	2.6	-1.8
urum	43.81	87.60	2.17	626	6.1 ± 1.4	30.4 ± 1.6	-6.9 ± 3.9	3.6	4.8	9.7	8.7	2.6
vill	40.44	356.05	6.11	2100	14.2 ± 0.4	19.1 ± 0.5	-2.8 ± 1.3	2.8	4.4	8.3	-0.7	-0.9
wsrt	52.91	6.60	3.51	1223	15.6 ± 0.5	17.3 ± 0.6	-1.5 ± 1.9	2.0	3.0	6.3	1.4	-1.3
yakz	62.03	129.68	3.13	945	-10.0 ± 1.8	22.8 ± 1.0	0.1 ± 2.6	6.8	4.5	9.2	1.8	3.2
zeck	43.79	41.57	3.26	1112	9.4 ± 0.6	26.9 ± 0.9	3.1 ± 2.3	2.5	4.1	8.1	0.5	0.9
India (In)												
HYDE	17.42	78.55	5.47	268	32.2 ± 0.8	42.1 ± 1.5	2.1 ± 2.5	4.4	9.5	15.2	0.3	0.3
IISC	13.02	77.57	5.98	1611	31.8 ± 0.6	42.4 ± 1.0	-0.1 ± 1.7	3.9	8.2	12.8	-0.1	-0.0
MALD <sup>l</sup>	4.19	73.53	1.39	485	30.1 ± 2.8	41.5 ± 4.2	8.6 ± 6.5	4.7	8.2	11.0	-1.8	-1.6
coco <sup>l</sup>	-12.19	96.83	4.55	1363	47.7 ± 0.7	43.0 ± 1.3	0.3 ± 2.2	3.9	8.5	12.4	17.9	3.0
dgar	-7.27	72.37	4.60	1556	29.8 ± 0.7	46.0 ± 1.4	2.5 ± 2.2	4.1	8.8	13.5	-2.0	2.8
Nazca (Nz)												
EISL	-27.15	250.62	6.94	1947	-9.0 ± 0.7	65.5 ± 1.0	1.8 ± 1.7	5.5	9.9	16.1	-0.2	-2.4
GALA	-0.74	269.70	4.91	1001	9.5 ± 0.7	53.0 ± 1.1	2.9 ± 1.9	4.0	7.4	10.6	1.3	1.9
GALD <sup>m</sup>	-0.90	-89.62	9	-	4.4 ± 5.8	56.3 ± 8.2	3.8 ± 7.0	-	-	-	-4.4	5.1
EASA <sup>m</sup>	-27.15	-109.38	12	-	-9.3 ± 0.7	68.1 ± 0.8	3.9 ± 0.8	-	-	-	-0.5	0.2
7097 <sup>k</sup>	-27.15	-109.38	17	-	-9.3 ± 0.8	68.1 ± 0.9	3.9 ± 0.9	-	-	-	-0.5	0.2
North America (Na)												
AMC2	38.80	-104.52	2.20	760	-8.9 ± 0.8	-11.5 ± 1.2	-4.2 ± 3.2	2.2	4.0	6.8	0.6	1.7
AML5	35.15	-101.88	4.93	1575	-8.1 ± 0.5	-12.3 ± 0.9	-5.4 ± 2.2	3.0	6.1	14.3	0.5	0.2
ANTO	29.49	-98.58	4.93	1661	-7.1 ± 0.5	-10.8 ± 0.9	-2.3 ± 2.4	3.2	6.5	17.0	0.3	0.2
AOML	25.73	-80.16	3.11	1051	-0.1 ± 0.7	-9.4 ± 1.4	-1.1 ± 2.7	2.5	6.2	9.7	0.3	1.0
ARL5	32.76	-97.06	4.93	1585	-6.3 ± 0.6	-13.4 ± 0.9	-5.5 ± 2.2	3.7	6.1	14.6	0.5	-1.3
ASHV	35.60	-82.55	3.16	1034	-0.6 ± 0.7	-13.9 ± 1.1	1.4 ± 2.6	2.7	5.1	9.1	0.7	-0.3
ATL1	33.94	-84.52	2.55	732	-2.1 ± 0.9	-13.9 ± 1.6	4.7 ± 3.7	2.8	5.8	11.5	0.0	-0.9
BARH	44.40	-68.22	2.25	807	3.2 ± 0.9	-14.8 ± 1.2	0.4 ± 3.3	2.3	4.0	7.8	-1.0	1.0
BRMU	32.37	-64.70	7.78	2697	6.1 ± 0.3	-12.4 ± 0.5	-0.4 ± 1.1	2.6	5.1	10.3	0.6	-0.2
BRU1	43.89	-69.95	4.97	1764	3.2 ± 0.5	-15.3 ± 0.7	0.2 ± 1.8	3.2	5.1	10.9	-0.3	0.5
Ccv1	28.46	-80.54	1.15	396	-3.6 ± 1.9	-15.9 ± 3.3	3.9 ± 8.0	2.5	5.3	11.1	-	-
Ccv3	28.46	-80.55	2.42	836	-0.9 ± 0.9	-12.6 ± 1.7	0.5 ± 3.6	2.7	5.9	10.2	-	-
CCV* <sup>i</sup>	28.46	-80.55	3.56	1232	-1.2 ± 0.6	-13.2 ± 1.1	0.9 ± 2.5	2.7	5.8	10.6	-0.6	-1.9
CHL1	38.78	-75.09	5.29	1806	0.7 ± 0.4	-15.4 ± 0.7	-0.7 ± 1.8	2.8	5.0	12.0	-0.8	-0.9
COLA	34.08	-81.12	2.00	712	-0.7 ± 1.0	-11.3 ± 1.8	-4.4 ± 4.0	2.4	5.3	8.8	0.1	1.8
COSA	33.57	-111.88	2.30	640	-12.4 ± 1.3	-12.1 ± 1.6	3.1 ± 3.6	3.5	5.2	8.9	-0.4	-1.1
DNRC	39.16	-75.52	2.77	856	0.7 ± 0.8	-13.9 ± 1.3	-3.0 ± 3.1	2.5	5.0	9.7	-0.7	0.7
DQUA	34.11	-94.29	3.97	1364	-5.3 ± 0.6	-13.0 ± 0.9	7.5 ± 2.5	3.0	5.2	13.0	0.5	-0.3
DUCK	36.18	-75.75	3.34	1054	1.3 ± 0.7	-14.4 ± 1.0	-8.7 ± 2.8	2.8	5.0	11.6	0.0	-0.7

Table 1. (continued)

Site <sup>a</sup>	Position <sup>b</sup>		$\Delta T$ , <sup>c</sup> years	Total Data <sup>d</sup>	Velocity <sup>e</sup>			WRMS <sup>f</sup>			Rate Res. <sup>g</sup>	
	°N	°E			North	East	Vertical	N	E	V	N	E
ERLA	39.02	-84.61	4.17	1263	-1.5 ± 0.6	-14.3 ± 0.7	0.4 ± 1.8	2.7	4.2	7.8	0.6	0.2
GAIT	39.13	-77.22	6.00	1705	0.6 ± 0.3	-14.0 ± 0.5	0.9 ± 1.3	2.3	4.5	8.6	-0.1	0.6
GALB	39.21	-84.49	2.99	999	-1.2 ± 0.7	-14.2 ± 1.1	-1.5 ± 2.7	2.5	4.6	9.2	0.9	0.4
GDAC	37.78	-102.18	3.99	1407	-7.2 ± 0.5	-14.0 ± 0.8	2.4 ± 2.0	2.4	4.8	8.9	1.5	-0.9
GLPT	37.25	-76.50	3.57	1227	1.3 ± 0.6	-14.4 ± 1.0	-3.9 ± 2.3	2.5	5.2	9.4	0.3	-0.3
GODE	39.02	-76.83	7.64	2642	0.9 ± 0.3	-14.0 ± 0.4	-0.9 ± 1.0	2.6	4.5	8.9	0.0	0.6
HBRK	38.30	-97.29	5.71	1938	-7.1 ± 0.4	-13.7 ± 0.6	0.0 ± 1.4	3.0	4.7	9.1	-0.2	0.0
HKLO	35.68	-95.86	5.71	1957	-6.3 ± 0.4	-13.4 ± 0.5	1.4 ± 1.5	3.0	4.5	9.7	0.1	-0.3
HVLK	37.65	-99.11	4.59	1587	-7.6 ± 0.5	-13.8 ± 0.7	0.3 ± 1.8	2.7	4.6	9.3	0.0	-0.4
KAN1	39.13	-95.40	4.30	1510	-5.4 ± 0.7	-13.8 ± 0.8	2.2 ± 2.0	3.7	5.1	10.2	0.8	0.3
KELY	66.99	-50.94	5.33	1693	10.2 ± 0.5	-18.3 ± 0.6	-3.7 ± 1.5	3.1	4.4	8.6	-0.2	0.0
KULU	65.58	-37.15	4.44	886	14.6 ± 0.6	-15.6 ± 0.6	0.6 ± 2.1	2.9	3.6	10.4	-0.2	-0.3
KYW1	24.58	-81.65	3.56	999	-0.6 ± 0.7	-9.7 ± 1.4	2.6 ± 2.6	3.1	7.1	11.2	0.4	0.3
LMNO	36.69	-97.48	5.86	2014	-7.0 ± 0.5	-13.3 ± 0.6	-3.4 ± 1.5	3.2	5.0	10.1	0.0	-0.1
LUBB	33.54	-101.84	4.93	1673	-9.0 ± 0.6	-12.2 ± 1.0	0.0 ± 2.3	3.6	6.9	16.4	-0.4	-0.2
MDO1	30.68	-104.01	7.57	2626	-9.7 ± 0.3	-11.3 ± 0.5	0.8 ± 1.0	2.7	5.0	8.3	-0.4	-0.3
Mia1	25.73	-80.16	2.82	921	-1.0 ± 0.8	-11.2 ± 1.5	3.1 ± 3.4	2.7	6.0	11.6	-	-
Mia3	25.73	-80.16	2.83	824	-0.5 ± 1.0	-11.3 ± 1.6	5.3 ± 3.2	3.2	6.3	10.9	-	-
MIA* <sup>i</sup>	25.73	-80.16	5.29	1745	-0.8 ± 0.5	-11.6 ± 0.8	4.1 ± 1.7	3.0	6.0	11.3	-0.4	-1.2
MNP1	41.07	-71.86	3.04	1031	3.5 ± 0.9	-15.9 ± 1.1	2.6 ± 2.7	3.2	5.0	9.4	0.7	-0.8
NDSK	37.38	-95.64	4.00	1386	-5.9 ± 0.6	-13.6 ± 0.8	1.3 ± 2.0	2.9	4.5	9.3	0.4	0.0
NLIB	41.77	-91.57	7.83	2744	-4.8 ± 0.3	-14.1 ± 0.4	-1.5 ± 1.0	2.6	4.3	8.4	0.0	0.9
ODS5	31.87	-102.32	4.93	1699	-9.5 ± 0.6	-11.8 ± 1.0	-1.7 ± 2.3	3.5	7.0	16.3	-0.8	-0.3
OMH1	41.78	-95.91	2.23	768	-6.0 ± 1.3	-13.8 ± 2.5	6.9 ± 4.2	3.4	7.8	11.7	0.4	1.0
PASO	31.77	-106.41	4.89	1638	-10.8 ± 0.5	-12.0 ± 0.8	-5.7 ± 2.0	3.2	5.7	13.1	-0.6	-0.9
PATT	31.78	-95.72	3.61	1227	-5.6 ± 0.6	-12.9 ± 1.2	3.4 ± 2.6	2.7	6.1	11.8	0.7	-1.0
PIT1	40.55	-79.70	3.87	1286	-1.4 ± 0.6	-13.8 ± 0.7	-8.6 ± 2.1	2.8	4.2	9.2	-1.2	1.2
PLTC	40.18	-104.73	5.86	2060	-10.4 ± 0.4	-13.4 ± 0.5	-2.3 ± 1.3	2.6	4.0	8.4	-0.8	0.1
Por2	43.07	-70.71	3.32	1112	2.5 ± 0.8	-16.1 ± 1.0	5.8 ± 2.7	3.1	5.0	10.7	-	-
Por4	43.07	-70.71	1.55	517	2.1 ± 1.8	-12.3 ± 2.7	3.0 ± 6.0	3.4	6.0	11.6	-	-
POR* <sup>i</sup>	43.07	-70.71	5.29	1629	2.5 ± 0.5	-15.6 ± 0.7	5.5 ± 1.7	3.2	5.6	11.0	-0.7	0.0
PRCO	34.98	-97.52	5.10	1722	-7.5 ± 0.4	-12.6 ± 0.7	-2.3 ± 1.6	2.7	4.9	9.7	-0.5	0.1
PRDS	50.87	-114.29	3.33	1090	-13.4 ± 0.6	-14.0 ± 0.8	-2.0 ± 2.1	2.3	3.7	7.0	-0.7	0.6
PSU1	40.81	-77.85	3.16	935	0.4 ± 0.7	-14.9 ± 1.0	-1.8 ± 2.8	2.5	4.4	10.1	-0.1	0.2
Rcm5	25.61	-80.38	3.05	862	0.5 ± 0.8	-9.2 ± 1.4	-3.7 ± 3.0	2.8	5.8	10.8	-	-
Rcm6	25.61	-80.38	1.89	623	-0.9 ± 1.0	-12.6 ± 2.2	6.4 ± 4.5	2.3	6.0	10.0	-	-
RCM* <sup>i</sup>	25.61	-80.38	4.94	1485	0.1 ± 0.5	-10.0 ± 0.9	-1.3 ± 1.9	2.6	6.1	11.5	0.6	0.4
RIC1	37.54	-77.43	3.72	1144	0.8 ± 0.6	-14.0 ± 0.8	-1.7 ± 2.3	2.4	4.2	9.6	0.2	0.2
SAV1	32.14	-81.70	2.09	643	-1.1 ± 1.1	-13.2 ± 1.6	1.9 ± 4.1	2.7	4.7	10.0	-0.1	-0.7
SOL1	38.32	-76.45	5.27	1391	1.2 ± 0.4	-14.1 ± 0.6	-4.8 ± 1.7	2.4	4.8	10.4	0.2	0.3
STJO	47.60	-52.68	8.00	2813	10.1 ± 0.3	-15.7 ± 0.4	-1.3 ± 1.0	2.8	4.8	8.4	0.3	-0.5
StI3	38.61	-89.76	3.93	1317	-4.1 ± 0.6	-13.7 ± 0.9	0.0 ± 2.3	2.6	5.2	11.2	-	-
StI4	38.61	-89.76	0.66	218	-11.9 ± 4.6	-14.9 ± 7.4	-24.0 ± 13.1	3.6	6.8	10.0	-	-
STL* <sup>i</sup>	38.61	-89.76	4.57	1535	-4.1 ± 0.5	-13.7 ± 0.8	0.0 ± 2.0	2.8	5.4	11.2	0.0	0.5
TCUN	35.09	-103.61	3.10	1017	-8.4 ± 0.7	-13.9 ± 0.9	5.0 ± 2.5	2.6	4.2	8.4	0.8	-1.6
THU1	76.54	-68.79	5.67	1803	3.5 ± 0.4	-22.0 ± 0.4	-0.7 ± 1.4	3.0	3.2	8.9	-0.5	-0.7
TMGO	40.13	-105.23	6.15	1882	-9.5 ± 1.4	-14.5 ± 0.6	-0.7 ± 1.3	10.6	5.1	8.1	0.3	-1.0
USNA	38.98	-76.48	5.27	1377	-0.5 ± 0.8	-14.1 ± 0.9	0.2 ± 1.5	4.9	6.7	8.5	-1.5	0.5
USNO	38.92	-77.07	3.67	1157	1.2 ± 0.6	-14.3 ± 0.8	-2.4 ± 2.2	2.4	4.4	8.7	0.4	0.3
VCIO	36.07	-99.22	5.71	1939	-6.3 ± 0.5	-12.6 ± 0.6	-0.6 ± 1.6	3.1	5.3	10.7	1.3	0.3
VIMS	37.61	-75.69	3.84	1207	1.4 ± 0.5	-14.5 ± 0.9	-2.3 ± 2.1	2.2	4.8	9.0	0.1	-0.3
WES2	42.61	-71.49	7.88	2721	2.3 ± 0.3	-14.6 ± 0.4	-2.3 ± 1.0	3.2	4.8	8.2	-0.6	0.9
WHN1	42.74	-103.33	2.17	759	-8.6 ± 0.9	-13.2 ± 1.3	4.7 ± 3.5	2.3	4.2	8.3	0.5	1.1
WLCI	40.81	-87.05	2.55	749	-4.3 ± 0.9	-13.7 ± 1.5	0.1 ± 3.3	2.6	5.3	9.6	-1.2	1.3
WSMN	32.41	-106.35	5.68	1939	-10.0 ± 0.5	-12.1 ± 0.6	0.8 ± 1.5	3.4	5.3	9.9	0.2	-0.8
algo	45.96	-78.07	8.00	2851	-0.6 ± 0.3	-16.3 ± 0.3	2.8 ± 0.9	2.7	3.9	7.5	-1.0	0.2
chb1	45.65	-84.47	5.29	1797	-1.3 ± 0.5	-14.7 ± 0.6	1.2 ± 1.5	2.9	4.7	8.8	0.8	1.6
chur	58.76	-94.09	7.68	2399	-6.8 ± 0.4	-18.2 ± 0.4	9.9 ± 1.0	3.2	4.3	8.1	-1.1	0.5
det1	42.30	-83.10	5.29	1847	-3.2 ± 0.6	-14.3 ± 0.9	0.5 ± 1.7	3.7	6.5	10.5	-1.7	1.2
dubo	50.26	-95.87	4.20	1449	-7.6 ± 0.5	-17.8 ± 0.6	1.8 ± 1.8	2.3	3.8	7.7	-1.2	-1.0
flin	54.73	-101.98	4.56	1603	-10.5 ± 0.5	-17.5 ± 0.6	3.7 ± 1.6	2.8	4.0	7.9	-1.9	-0.4
kew1	47.23	-88.62	3.56	1252	-5.1 ± 0.7	-17.6 ± 0.9	2.4 ± 2.3	2.9	4.8	9.2	-1.4	-1.0
mil1	43.00	-87.89	3.56	1247	-3.0 ± 0.7	-17.3 ± 0.9	0.2 ± 2.3	2.8	4.6	9.4	0.5	-1.8
neb3	46.32	-84.15	3.66	1277	1.1 ± 0.6	-15.1 ± 0.9	4.4 ± 2.2	2.8	4.6	8.9	3.1	1.4
nrc1	45.45	-75.62	6.59	1939	0.7 ± 0.3	-15.8 ± 0.4	2.7 ± 1.2	2.4	4.2	8.3	-0.6	0.5
oro_sag1	44.90	-68.67	2.33	615	4.2 ± 0.9	-16.3 ± 1.2	-0.9 ± 3.4	2.4	4.1	8.3	0.2	-0.3
sag1	43.63	-83.84	4.97	1715	-2.8 ± 0.5	-18.4 ± 0.7	1.2 ± 1.7	2.9	5.2	10.0	-1.0	-2.6
sch2	54.83	-66.83	3.33	1150	5.2 ± 0.7	-17.8 ± 0.8	9.9 ± 2.3	2.8	4.1	8.5	0.5	0.4
stb1	44.80	-87.31	3.56	1221	-2.5 ± 0.8	-16.8 ± 1.0	2.0 ± 2.2	3.2	5.0	8.4	0.7	-0.8
stp1	44.30	-91.90	3.56	1251	-4.4 ± 0.6	-16.0 ± 1.0	-2.2 ± 2.4	2.6	5.0	9.8	0.5	-0.3
vcap	44.26	-72.58	4.57	1402	2.2 ± 0.6	-16.0 ± 0.8	1.4 ± 2.0	3.3	5.0	10.8	-0.3	-0.1
whp1	46.77	-84.96	4.02	1335	-2.7 ± 0.6	-16.3 ± 0.7	3.2 ± 2.1	2.8	4.2	9.4	-0.4	0.3

**Table 1.** (continued)

Site <sup>a</sup>	Position <sup>b</sup>		$\Delta T^c$ years	Total Data <sup>d</sup>	Velocity <sup>e</sup>			WRMS <sup>f</sup>			Rate Res. <sup>g</sup>	
	°N	°E			North	East	Vertical	N	E	V	N	E
wis1	46.71	-92.02	3.56	1246	-4.9 ± 0.6	-17.5 ± 0.8	-1.3 ± 2.3	2.5	4.1	9.0	0.1	-1.2
yell	62.48	-114.48	8.00	2840	-13.4 ± 0.3	-16.4 ± 0.4	2.6 ± 1.0	3.3	4.2	9.3	-0.6	0.0
you1	43.23	-78.97	3.56	1254	-1.8 ± 0.7	-14.2 ± 1.1	0.9 ± 2.6	3.1	5.6	11.5	-1.8	1.6
bili	68.08	-193.56	1.33	369	-21.1 ± 2.0	11.4 ± 1.6	1.2 ± 5.9	3.1	3.1	8.1	-1.0	3.4
chal	32.76	-79.84	5.29	1824	-0.7 ± 0.5	-10.8 ± 0.8	2.2 ± 1.7	2.9	6.1	10.5	-0.4	1.9
drao	49.32	-119.62	8.00	2848	-13.5 ± 0.3	-12.4 ± 0.4	-0.4 ± 1.0	3.0	4.6	8.1	0.9	0.9
fair	64.98	-147.50	8.00	2739	-22.7 ± 0.4	-7.4 ± 0.4	-2.3 ± 1.2	4.0	4.9	11.9	-2.2	0.4
hofn	64.27	-15.20	3.24	1093	14.7 ± 0.7	11.4 ± 0.8	4.0 ± 2.3	2.7	3.8	7.7	-5.1	20.6
mem2	35.47	-90.21	5.29	1779	-3.6 ± 0.5	-15.7 ± 0.8	0.8 ± 1.7	3.1	5.8	11.0	0.7	-2.4
pie1	34.30	-108.12	8.00	2826	-11.6 ± 0.3	-11.7 ± 0.4	1.6 ± 0.9	2.7	4.7	8.0	-0.9	-0.1
reyk	64.14	-21.96	4.95	1676	18.8 ± 0.6	-11.4 ± 0.6	-3.4 ± 1.5	3.4	4.4	7.9	0.2	-0.2
Nubia (Nu)												
GOUG	-40.35	350.12	2.42	653	15.3 ± 2.0	19.1 ± 2.3	-2.6 ± 4.3	5.7	7.6	13.1	-0.9	-1.6
Hart	-25.89	27.71	4.44	1427	18.3 ± 1.5	16.3 ± 2.4	-0.3 ± 2.9	8.0	14.5	18.8	-	-
Hark	-25.89	27.71	3.15	997	16.0 ± 1.2	17.9 ± 1.6	9.8 ± 3.2	4.4	7.0	12.9	-	-
Harb	-25.89	27.71	0.62	102	20.7 ± 6.3	11.3 ± 7.7	2.9 ± 16.1	4.3	6.1	11.0	-	-
HAR* <sup>i</sup>	-25.89	27.71	8.00	2528	17.4 ± 0.7	17.0 ± 1.0	3.7 ± 1.5	6.3	10.9	16.4	1.0	-0.7
HRAO	-25.89	27.69	4.26	1099	15.5 ± 0.8	18.3 ± 1.3	0.8 ± 2.2	4.2	7.7	10.9	-0.9	0.6
Mas1 <sup>h</sup>	27.76	344.37	6.58	2116	15.6 ± 0.4	16.2 ± 0.6	0.6 ± 1.3	2.8	5.8	10.0	-	-
Masp <sup>h</sup>	27.76	344.37	1.69	600	13.1 ± 1.8	14.6 ± 3.0	-8.1 ± 5.2	3.6	7.3	10.8	-	-
MAS* <sup>n</sup>	27.76	344.37	8.00	2443	15.5 ± 0.3	16.4 ± 0.5	1.0 ± 1.1	3.0	6.0	10.4	0.0	0.1
SUTH	-32.38	20.81	2.68	891	15.9 ± 1.4	17.6 ± 1.7	-0.5 ± 3.3	4.4	6.6	10.8	-1.0	0.6
katz	33.00	35.69	4.18	590	19.3 ± 0.7	22.7 ± 1.0	3.9 ± 2.4	3.2	5.3	11.4	3.8	0.0
lamp	35.50	12.61	1.77	586	17.5 ± 1.3	19.5 ± 1.4	4.5 ± 4.4	2.7	3.7	8.5	0.3	0.9
nico	35.14	33.40	3.36	959	13.6 ± 0.8	19.0 ± 0.9	1.1 ± 2.3	3.0	4.2	8.0	-2.2	-3.1
Iavh <sup>h</sup>	33.98	353.13	0.59	170	14.7 ± 3.0	24.0 ± 5.1	-4.3 ± 12.1	2.0	4.2	6.7	-	-
Rabt <sup>h</sup>	33.98	353.14	0.56	133	18.7 ± 7.4	40.9 ± 9.7	-43.2 ± 17.6	4.8	7.2	11.6	-	-
rab* <sup>n</sup>	33.98	353.14	2.54	303	18.9 ± 1.1	18.1 ± 1.7	-7.0 ± 3.5	3.0	5.5	8.7	2.4	2.4
mali	-3.00	40.19	5.12	1714	12.7 ± 0.7	26.8 ± 1.4	1.5 ± 2.3	4.3	10.1	16.3	-2.2	5.1
sey*	-4.67	55.48	3.98	276	8.0 ± 1.3	26.3 ± 2.8	1.9 ± 3.7	5.1	12.8	17.8	-4.0	5.1
Okhotsk (Ok)												
GS05	42.15	139.52	4.67	280	-12.3 ± 0.8	0.8 ± 1.0	-11.6 ± 2.6	3.7	5.6	12.9	-0.9	-1.4
STKW	43.53	141.85	4.46	1257	-13.3 ± 0.9	3.7 ± 1.0	-2.0 ± 3.1	4.9	6.4	20.2	-0.5	0.8
TKBA	36.11	140.09	4.46	1031	-10.1 ± 0.8	-4.2 ± 1.0	-6.3 ± 2.7	4.1	6.0	16.0	1.6	-1.4
TSKB	36.11	140.09	7.00	2445	-12.7 ± 0.7	-4.2 ± 0.7	-2.3 ± 1.3	5.8	6.9	11.7	-1.0	-1.4
USUD	36.13	138.36	8.00	2716	-10.1 ± 0.6	0.0 ± 0.7	0.2 ± 1.3	5.8	8.1	13.4	0.6	2.6
mag0	59.58	150.77	3.13	908	-23.8 ± 1.1	14.9 ± 0.9	-1.4 ± 2.4	4.2	4.3	7.6	-5.7	1.3
petp	53.07	158.61	2.23	694	-9.50 ± 1.6	0.2 ± 1.3	3.9 ± 3.8	3.4	4.2	9.5	12.9	-6.1
Pacific (Pa)												
CHAT	-43.96	183.43	5.24	1879	31.3 ± 0.6	-39.7 ± 0.8	3.6 ± 1.5	3.6	5.7	9.2	1.4	0.4
FALE	-13.83	188.00	2.00	573	30.5 ± 1.7	-63.2 ± 3.5	-3.2 ± 5.1	4.1	9.5	13.0	-0.2	-1.3
HNLC	21.30	202.14	2.00	373	32.4 ± 1.6	-61.2 ± 2.7	-1.1 ± 5.2	3.7	7.1	12.3	0.7	0.2
KOK1	21.98	200.24	3.56	1235	32.0 ± 0.8	-61.8 ± 1.4	2.6 ± 2.8	3.6	7.1	13.4	0.3	-0.3
KOKB	22.13	200.34	7.99	2768	30.8 ± 0.5	-60.8 ± 0.7	1.6 ± 1.3	4.5	7.7	13.3	-0.9	0.6
KWJ1	8.72	167.73	4.79	1442	25.6 ± 0.8	-67.9 ± 1.2	0.0 ± 2.3	4.4	8.1	14.6	-0.4	-0.2
MARC	24.29	153.98	4.83	1102	19.4 ± 1.0	-71.2 ± 1.2	5.6 ± 2.4	5.4	7.8	15.2	-1.5	-1.5
Pama <sup>h</sup>	-17.57	210.43	4.24	1423	31.5 ± 1.4	-63.1 ± 3.1	1.7 ± 3.5	7.1	18.2	23.5	-	-
Tahi <sup>h</sup>	-17.58	210.39	1.02	254	33.6 ± 3.7	-62.1 ± 6.8	1.2 ± 11.3	4.4	9.2	15.1	-	-
Thi <sup>h</sup>	-17.58	210.39	2.56	618	31.2 ± 1.5	-63.3 ± 2.7	8.9 ± 4.5	4.4	9.4	15.2	-	-
THT* <sup>i</sup>	-17.58	210.39	7.98	2295	31.4 ± 0.7	-63.3 ± 1.4	3.6 ± 1.7	6.3	14.8	20.0	-0.1	0.6
TRUK	7.45	151.89	3.88	763	20.7 ± 0.8	-69.7 ± 1.4	5.8 ± 2.8	3.5	7.3	13.5	0.7	-1.4
Cice <sup>h</sup>	31.87	243.33	3.90	1316	15.5 ± 0.6	-39.1 ± 0.9	-0.7 ± 1.9	2.6	4.9	7.9	-	-
Cic1 <sup>h</sup>	31.87	243.33	1.76	613	20.5 ± 1.2	-39.2 ± 1.8	4.1 ± 4.3	2.6	4.5	7.9	-	-
cic* <sup>i</sup>	31.87	243.33	5.76	1929	15.9 ± 0.4	-39.1 ± 0.6	-0.3 ± 1.3	3.0	4.8	8.2	-8.2	5.9
farb	37.70	237.00	4.26	1364	22.3 ± 0.5	-39.6 ± 0.7	0.0 ± 1.9	2.4	4.5	9.2	-3.9	1.6
guam	13.59	144.87	5.94	2044	0.6 ± 0.6	-8.9 ± 1.0	5.2 ± 1.8	4.3	8.7	14.9	-16.3	61.3
mkea	19.80	204.54	4.25	1394	31.2 ± 0.7	-61.9 ± 1.0	1.1 ± 2.1	3.3	6.3	10.5	-0.5	-0.4
pal* <sup>i,1</sup>	7.34	134.48	4.28	563	19.0 ± 0.8	-64.2 ± 1.8	-4.8 ± 2.8	3.8	9.6	14.4	7.3	4.7
scip	32.91	241.51	2.94	723	20.4 ± 0.7	-42.1 ± 1.1	0.7 ± 2.5	2.2	4.6	6.7	-4.3	2.4
smi1	33.25	240.48	5.11	1148	20.4 ± 0.4	-41.7 ± 0.7	-0.3 ± 1.6	2.6	4.7	8.4	-4.7	2.7
spm <sup>x</sup>	31.05	244.53	2.19	386	20.4 ± 1.0	-46.7 ± 1.6	2.4 ± 3.6	2.4	4.8	7.6	-3.3	-1.2
upo1	20.25	204.12	3.56	1242	34.6 ± 1.2	-62.9 ± 1.7	-4.3 ± 3.0	5.0	8.5	14.4	2.9	-1.5
vndp	34.56	239.38	7.99	2557	19.3 ± 0.4	-42.2 ± 0.5	-1.0 ± 1.1	3.9	5.7	10.5	-6.1	1.3
Philippine (Ph)												
GS11	25.83	131.23	4.67	1147	21.4 ± 0.8	-38.1 ± 0.9	2.4 ± 2.4	4.1	6.1	15.2	-0.5	-0.6
GS13	25.95	131.29	3.58	1045	22.5 ± 0.9	-36.4 ± 1.3	-3.2 ± 2.9	4.0	6.3	13.9	0.6	1.0
OKTO	20.43	136.08	3.91	14	14.0 ± 1.5	-45.7 ± 2.6	9.4 ± 5.3	4.0	7.5	15.2	-2.3	-0.3
Pala <sup>h</sup>	7.34	134.48	0.66	178	15.8 ± 5.1	-50.6 ± 11.1	1.0 ± 17.2	3.9	9.8	14.8	-	-

Table 1. (continued)

Site <sup>a</sup>	Position <sup>b</sup>		$\Delta T$ , <sup>c</sup> years	Total Data <sup>d</sup>	Velocity <sup>e</sup>			WRMS <sup>f</sup>			Rate Res. <sup>g</sup>	
	$^{\circ}$ N	$^{\circ}$ E			North	East	Vertical	N	E	V	N	E
Pbls <sup>h,i</sup>	7.34	134.48	2.53	385	19.3 ± 1.3	-65.0 ± 2.9	-7.1 ± 8.8	3.6	9.3	34.1	-	-
PAL* <sup>i,l</sup>	7.34	134.48	4.28	563	19.0 ± 0.8	-64.2 ± 1.8	-4.8 ± 2.8	3.8	9.6	14.4	0.8	0.2
ccjm	27.10	142.18	4.29	1167	9.3 ± 1.0	-39.2 ± 1.5	-0.8 ± 2.7	5.2	8.7	15.8	0.3	-5.0
gsi2	26.64	142.16	3.75	1058	6.6 ± 1.2	-33.1 ± 1.7	-2.8 ± 3.1	5.3	9.0	15.7	-2.4	1.9
guam	13.59	144.87	5.94	2044	0.6 ± 0.6	-8.9 ± 1.0	5.2 ± 1.8	4.3	8.7	14.9	-5.1	46.3
haci	33.07	139.82	4.43	1057	6.3 ± 0.8	-14.8 ± 1.2	-1.0 ± 2.3	4.2	7.0	12.3	-5.5	9.6
oknw	26.14	127.77	4.5	1278	-39.7 ± 0.9	35.3 ± 1.2	0.4 ± 2.3	5.0	7.7	13.1	-65.6	-73.0
sl02	22.04	121.56	5.00	328	41.7 ± 1.5	-21.0 ± 2.1	24.8 ± 3.4	7.7	12.2	23.6	9.0	24.2
South America (Sa)												
ASC1	-7.95	345.59	4.66	1542	8.5 ± 0.8	-5.1 ± 1.3	0.3 ± 2.2	4.3	8.5	13.4	-0.6	0.8
BOMJ	-13.26	316.58	2.92	661	10.1 ± 1.2	-5.7 ± 2.1	0.3 ± 3.4	4.0	8.3	11.8	-0.6	-0.6
BRAZ	-15.95	312.12	5.83	1434	10.5 ± 0.6	-5.5 ± 1.0	2.9 ± 1.8	4.0	7.6	13.1	-0.2	-0.7
FORT	-3.88	321.57	7.59	2634	11.5 ± 0.5	-6.9 ± 0.8	0.2 ± 1.5	4.4	8.7	16.1	0.9	-1.7
IMPZ	-5.49	312.50	2.67	518	10.6 ± 1.2	-9.1 ± 2.5	-3.4 ± 3.7	3.7	8.9	11.7	-0.1	-4.0
KOUR	5.25	307.19	7.99	2750	10.0 ± 0.5	-3.5 ± 0.8	1.7 ± 1.4	4.4	8.5	15.8	-0.6	1.8
LPGS	-34.91	302.07	5.50	1607	10.8 ± 0.8	-1.6 ± 1.0	5.5 ± 1.9	5.3	8.0	14.2	0.4	1.3
MANU	-3.12	299.94	3.06	617	8.7 ± 1.3	-4.2 ± 2.2	-2.6 ± 6.1	4.6	9.0	29.2	-1.6	0.8
PARA	-25.45	310.77	3.18	834	11.3 ± 1.3	-4.8 ± 1.6	7.0 ± 3.1	4.8	7.2	12.1	0.7	-0.4
UEPP	-22.12	308.59	3.40	947	11.4 ± 1.2	-4.9 ± 1.6	6.0 ± 3.0	4.8	7.4	13.4	0.8	-0.5
VICO	-20.76	317.13	2.80	737	10.2 ± 1.3	-6.5 ± 1.9	8.1 ± 3.6	4.3	7.3	12.7	-0.5	-1.5
South China (Cs)												
S01R	23.65	119.59	5.00	205	-17.8 ± 0.9	35.1 ± 1.6	9.1 ± 3.4	4.4	9.0	19.9	-1.3	1.1
SHAO	31.10	121.20	5.96	1827	-16.7 ± 0.5	34.0 ± 0.8	-0.1 ± 1.5	3.6	6.7	10.1	0.2	0.4
WUHN	30.53	114.36	4.93	1642	-14.8 ± 0.6	33.9 ± 0.8	0.2 ± 1.8	3.8	5.9	10.8	0.3	-0.7
bjfs	39.61	115.89	1.20	419	-13.8 ± 2.4	29.9 ± 2.8	-3.9 ± 7.5	3.4	4.9	11.0	1.7	-3.7
kunm	25.03	102.80	2.23	684	-21.2 ± 1.4	32.3 ± 2.5	-4.3 ± 4.5	3.7	7.8	13.1	-9.6	-3.5
oknw	26.14	127.77	4.5	1278	-39.7 ± 0.9	35.3 ± 1.2	0.4 ± 2.3	5.0	7.7	13.1	-21.3	2.4
urum	43.81	87.60	2.17	626	6.1 ± 1.4	30.4 ± 1.6	-6.9 ± 3.9	3.6	4.8	9.7	12.4	-6.4
xian	34.37	109.22	3.48	991	-14.7 ± 0.9	34.5 ± 1.1	0.1 ± 2.5	3.8	5.2	10.0	-1.1	-0.6
Sierra (Sr)												
CMBB	38.03	239.61	7.06	2133	-7.1 ± 0.4	-20.1 ± 0.6	3.2 ± 1.9	3.7	6.2	19.4	-0.1	2.0
MUSB	37.17	240.69	3.15	804	-6.8 ± 0.7	-24.2 ± 1.1	3.6 ± 2.6	2.4	4.7	8.5	-0.7	-2.7
ORVB	39.55	238.50	4.00	1225	-8.5 ± 0.5	-23.7 ± 0.7	-1.8 ± 2.1	2.2	4.2	9.5	-0.5	-0.5
SUTB	39.21	238.18	3.70	1061	-8.1 ± 0.6	-23.0 ± 0.8	-2.7 ± 2.1	2.4	4.1	7.7	0.1	-0.1
UCD1	38.54	238.25	3.82	896	-7.3 ± 0.6	-23.5 ± 0.8	-8.8 ± 2.4	2.6	4.2	10.3	0.9	-1.1
CEDA	35.75	-118.59	5.12	26	-5.1 ± 1.0	-21.0 ± 1.5	-2.7 ± 2.4	3.6	6.2	7.7	-0.5	0.4
KMED	36.02	-118.14	6.90	29	-4.5 ± 0.7	-20.2 ± 0.8	-1.3 ± 2.0	3.2	4.8	9.0	0.5	0.7
SPRN	36.18	-118.73	5.08	27	-7.9 ± 1.1	-18.0 ± 1.6	-4.5 ± 3.3	3.9	6.9	13.1	-3.2	-0.8
TIOG	37.93	-119.25	4.00	11	-6.4 ± 0.9	-24.0 ± 2.1	-1.1 ± 3.5	2.2	6.1	8.1	1.3	0.3
quin	39.97	239.06	7.99	2371	-8.0 ± 0.4	-19.8 ± 0.5	-2.6 ± 1.2	3.7	5.5	11.2	-0.5	3.7
Somalia (So)												
MALI	-3.00	40.19	5.12	1714	12.7 ± 0.7	26.8 ± 1.4	1.5 ± 2.3	4.3	10.1	16.3	0.0	-0.1
Sey1 <sup>h</sup>	-4.67	55.48	5.63	360	9.9 ± 1.0	33.0 ± 2.4	10.6 ± 3.1	5.9	15.6	22.2	-	-
SEY* <sup>o</sup>	-4.67	55.48	3.98	276	8.0 ± 1.3	26.3 ± 2.8	1.9 ± 3.7	5.1	12.8	17.8	0.0	0.2
Sunda (Su)												
BAKO	-6.49	106.85	2.83	765	-8.5 ± 1.3	24.0 ± 2.1	-4.6 ± 3.9	4.2	8.3	14.6	-0.4	0.5
NTUS	1.35	103.68	3.27	841	-6.0 ± 1.1	27.8 ± 2.1	-1.8 ± 3.2	4.1	9.2	13.5	0.3	-0.4
kunm	25.03	102.80	2.23	684	-21.2 ± 1.4	32.3 ± 2.5	-4.3 ± 4.5	3.7	7.8	13.1	-6.7	4.8

<sup>a</sup>Upper case sites are in stable plate interior used to calculate angular velocity for that plate; lowercase sites are in boundary zone or other deforming zone, sites with asterisk indicate composite site (see text).

<sup>b</sup>Site location in decimal degrees.

<sup>c</sup>Length of time series.

<sup>d</sup>Number of station days.

<sup>e</sup>Velocities in mm/yr relative to ITRF-97,  $\pm 1$  standard error.

<sup>f</sup>Weighted root mean square scatter of the daily position estimates about a best fit straight line.

<sup>g</sup>Residual rate with respect to stable plate angular velocity.

<sup>h</sup>Site that is part of a composite site solution (see footnotes i and n).

<sup>i</sup>Composite site solution that is made up of two sites linked with an offset (see text).

<sup>j</sup>Site Ankr without data after 15 August 1999 to avoid coseismic and postseismic effects of  $M > 7$  earthquakes.

<sup>k</sup>SLR site.

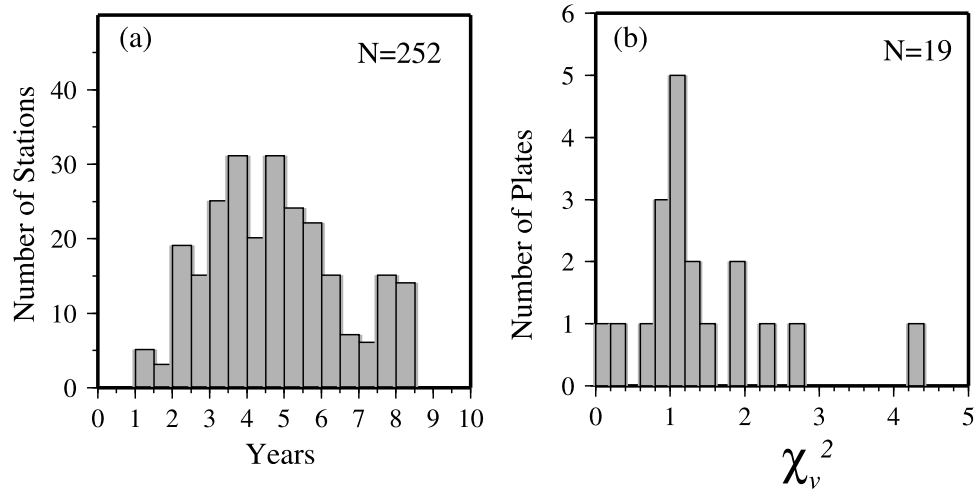
<sup>l</sup>Site for which an offset has been calculated even though no change in antenna type, height, or dome change was reported.

<sup>m</sup>DORIS site (GALD is really GALB).

<sup>n</sup>A published site tie was used to link two sites together.

<sup>o</sup>Site Sey1 without 1995 and 1996 data (see text).





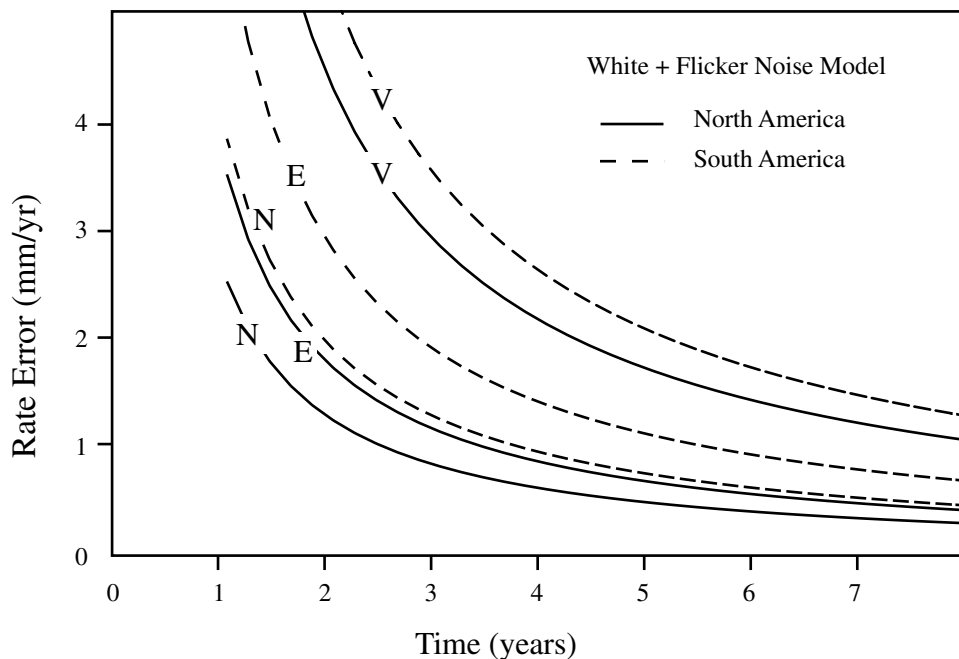
**Figure 2.** (a) Histogram showing length of time series for the sites analyzed in this study. (b) Histogram of  $\chi_v^2$  ( $\chi^2$  per degree of freedom) calculated for the 19 plates and blocks described in this study.

[9] The total time span of observations has an important influence on the uncertainty of site velocities from space geodetic data [Coates *et al.*, 1985; Dixon, 1991; Zhang *et al.*, 1997; Mao *et al.*, 1999]. Table 1 and Figure 2a show the time span of data for the sites used in this study. The great majority of sites in this study have recorded more than 300 days per year for more than 3.0 years, the mean being 4.7 years. Figure 3 shows the expected evolution of velocity error as a function of time for a typical station that operates essentially continuously, recording 300 days per year. By reference to Figure 3, we can see that on average, the velocity errors for the horizontal components of a 4-year time series will lie in the range from 0.5 to 1.5 mm/yr. However, in order to obtain adequate geographic coverage we have also used stations with time spans as short as 1.4 years. This type of station will have horizontal

component velocity uncertainties of  $\sim 2\text{--}5$  mm/yr with the analytical techniques used here.

[10] Table 1 lists the available data per plate. Eight plates or blocks (Amuria, Anatolia, Arabia, India, Philippine, Somalia, South China, and Sunda) have less than 10 velocity data (i.e., less than five sites, with two horizontal velocity components per site); hence their angular velocities may not be well determined.

[11] From the site velocity data listed in Table 1 we seek the best fit angular velocity vector (Euler vector) describing the relative motion of adjacent plates (the relative velocities of nonadjacent plates can also be derived, but the results are more difficult to test against geologic data). We first select sites thought to be representative of stable plate interiors, using standard geologic and seismological criteria. For example, for the Eurasian



**Figure 3.** Estimate of GPS velocity error as a function of time this study, calculated according to the model of Mao *et al.* [1999] as modified by Dixon *et al.* [2000a]. Solid lines show north (N), east (E) and vertical (V) error for a typical North American site, assuming WRMS for each component is 3, 5, and 10 mm, close to the mean of the 64 sites listed in Table 1. Dashed lines show corresponding curves based on mean South American site.

**Table 2.** Plate Angular Velocity Relative to ITRF–97 in Geographic Coordinates

Plate <sup>a</sup>	$\lambda$ , <sup>b</sup> °N	$\phi$ , <sup>b</sup> °E	$\omega$ , <sup>b</sup> deg/Myr	$\sigma_{\text{maj}}$ <sup>c</sup>	$\sigma_{\text{min}}$ <sup>c</sup>	$\zeta$ , <sup>d</sup> deg	$\sigma_{\omega}$ , deg/Myr	Number of Sites <sup>e</sup>	$\chi^2_{\nu}$ <sup>f</sup>	RMS Res. <sup>g</sup>		Mean <sup>h</sup> Rate Res.
										N	E	
Am	63.75	-133.76	0.327	23.5	1.6	-64	0.057	3	2.30	1.2	0.2	1.1
An	58.48	-134.00	0.226	1.6	1.0	32	0.010	7	1.02	0.9	1.3	1.0
Ar	51.47	2.89	0.521	3.1	0.7	-70	0.024	3	0.85	2.6	2.6	2.6
At	42.19	26.18	0.882	8.2	0.7	-65	0.889	4	1.60	3.7	4.5	4.1
Au	34.86	38.26	0.627	1.2	0.4	-65	0.004	11	2.78	0.6	1.4	1.3
Ca	37.74	-90.86	0.272	3.9	1.1	-50	0.021	6	0.91	0.9	1.0	1.2
Cs	54.58	-109.21	0.340	16.6	1.0	-40	0.007	3	1.27	0.8	0.8	1.0
Eu	58.27	-102.21	0.257	1.5	0.4	34	0.003	15	1.02	0.6	0.5	0.7
In	53.65	-13.99	0.483	11.7	0.5	80	0.013	3	0.26	1.0	1.0	1.0
Na	-2.39	-79.08	0.199	0.8	0.3	-6	0.002	64	1.05	0.6	0.8	0.9
Nu	52.25	-80.18	0.253	1.6	0.9	-83	0.004	5	0.82	0.8	0.9	1.1
Nz	44.45	-99.49	0.647	2.9	0.6	5	0.011	5	1.96	2.1	2.7	2.5
Ok	-37.92	-58.14	0.428	6.1	1.4	-85	0.124	5	4.36	1.0	1.6	1.8
Pa	-64.21	112.74	0.655	0.7	0.4	75	0.004	9	1.20	0.8	0.9	1.1
Ph	-46.52	-30.41	0.910	2.2	0.4	-35	0.046	4	1.03	1.3	0.6	1.3
Sa	-25.83	-135.38	0.106	7.7	2.8	-80	0.003	11	1.31	0.7	1.6	1.5
So	53.51	-101.55	0.310	0.9	0.1	42	0.002	2	0.01	0.0	0.1	0.1
Sr	10.87	-112.00	0.445	9.8	0.5	-14	0.099	9	1.96	1.2	1.2	1.5
Su	38.86	-86.94	0.393	10.2	0.8	-20	0.062	2	0.24	0.3	0.4	0.5

<sup>a</sup>Plate abbreviations are listed in Table 1.

<sup>b</sup>Parameter  $\lambda$  is latitude and  $\phi$  is longitude of the rotation pole in decimal degrees;  $\omega$  is the rotation rate.

<sup>c</sup>Two-dimensional 1-sigma lengths in degrees of the semi-major  $\sigma_{\text{maj}}$  and semi-minor axes  $\sigma_{\text{min}}$  of the pole error ellipse.

<sup>d</sup>The parameter  $\zeta$  is the azimuth of the semi-major ellipse axis in degrees clockwise from north.

<sup>e</sup>List of sites used is given in Table 1.

<sup>f</sup>The parameter  $\chi^2$  per degree of freedom.

<sup>g</sup>Root mean square of the rate residual north and east components in mm/yr.

<sup>h</sup>Mean rate residual in mm/yr.

plate we omit all data likely to be influenced by deformation associated with the northward motion of Africa, Arabia, or India, as indicated by seismicity, and also all sites east of longitude 130°E because they may lie within the deforming zone between Eurasia and North America or lie on the North American plate. In most cases we also choose sites located >100 km from significant plate boundary zone-related seismicity. In this way we hope to avoid seismic cycle effects, including interseismic strain accumulation, which can extend far from faults, as well as postseismic effects, which can persist long after an earthquake because of

viscous relaxation in the lower crust and upper mantle [e.g., *Pollitz, 1992; Pollitz and Sacks, 1994*]. By avoiding these effects our measured decadal velocities should be similar to longer-term plate velocities. In cases where we have had to violate this criterion in order to obtain sufficient data we have tested sensitivity to first-order strain effects via simple elastic half-space models.

[12] Next we invert the site velocity data to derive best fit angular velocities for each plate, termed here ITRF-97 angular velocities (Tables 2 and 3), minimizing the weighted, least squares

**Table 3.** Plate Angular Velocities Relative to ITRF-97 in Cartesian Coordinates With Covariance Matrix

Plate <sup>a</sup>	$\omega X$ <sup>b</sup>	$\omega Y$ <sup>b</sup>	$\omega Z$ <sup>b</sup>	$xx$ <sup>c</sup>	$xy$ <sup>c</sup>	$xz$ <sup>c</sup>	$yy$ <sup>c</sup>	$yz$ <sup>c</sup>	$zz$ <sup>c</sup>
Am	-1.746969	-1.824311	5.121903	0.857324	-1.063607	-1.096796	1.347375	1.378061	1.438827
An	-1.431031	-1.481635	3.358330	0.002206	0.000481	-0.001042	0.005049	-0.008574	0.028016
Ar	5.653346	0.285876	7.109457	0.110933	0.107269	0.082829	0.111556	0.083046	0.072361
At	10.231864	5.029467	10.335698	101.943568	65.724799	100.431138	42.396490	64.759834	98.973900
Au	7.046423	5.557522	6.250947	0.009488	-0.009391	0.007173	0.012160	-0.008208	0.009718
Ca	-0.056119	-3.751012	2.903333	0.029068	-0.061974	0.019386	0.144467	-0.044556	0.020405
Cs	-1.131332	-3.247545	4.836032	0.253516	-0.454367	-0.297816	0.831822	0.542589	0.366409
Eu	-0.499530	-2.309197	3.821139	0.002968	0.000972	0.004096	0.000727	0.001536	0.007045
In	4.852852	-1.208854	6.795223	0.064093	0.289369	0.072830	1.335119	0.335702	0.088999
Na	0.658291	-3.411026	-0.144969	0.000114	-0.000075	0.000066	0.001038	-0.000799	0.000962
Nu	0.459991	-2.658697	3.484752	0.007755	-0.000608	0.001160	0.001553	-0.000336	0.004719
Nz	-1.329904	-7.952151	7.910255	0.015252	0.034933	0.015391	0.140454	0.058358	0.036803
Ok	3.107407	-5.000841	-4.587389	1.822874	-1.538042	-1.838730	1.317255	1.561148	1.884001
Pa	-1.923772	4.589783	-10.300427	0.009299	0.001117	-0.000448	0.002121	-0.000067	0.005219
Ph	9.423451	-5.530138	-11.523673	0.316389	-0.337266	-0.192008	0.370821	0.208181	0.131048
Sa	-1.189878	-1.174240	-0.809045	0.015218	-0.013617	-0.003420	0.015430	0.003466	0.005410
So	-0.644017	-3.150603	4.346639	0.002358	0.002195	-0.000185	0.002151	-0.000177	0.000253
Sr	-2.860888	-7.064668	1.463571	0.615256	1.043353	-0.941995	1.777408	-1.602858	1.452819
Su	0.285162	-5.332412	4.302467	0.128464	-0.464367	0.020302	1.736351	-0.075708	0.016525

<sup>a</sup>Plate abbreviations are listed in Table 1.

<sup>b</sup>Units for  $\omega$  are  $10^{-3}$  rads/Myr with the X, Y, Z axes being parallel to (0°N,0°E), (0°N,90°E), and 90°N, respectively.

<sup>c</sup>These six columns give the indicated elements of the covariance matrix ( $10^{-6}$  rads<sup>2</sup>/Myr<sup>2</sup>).

**Table 4.** Relative Angular Velocities for Adjacent Plate Pairs

Plate Pair <sup>a</sup>	$\lambda^b$ , °N	$\phi^b$ , °E	$\omega$ , deg/Myr	$\sigma_{\text{maj}}^c$	$\sigma_{\text{min}}^c$	$\zeta$ , <sup>d</sup> deg	$\sigma_{\omega}$ , deg/ Myr
Am-Cs	10.45	113.39	0.090	56.7	7.5	21	0.113
Am-Eu	44.18	158.76	0.107	33.3	6.6	88	0.100
Am-Ok	59.14	146.80	0.648	8.1	2.0	25	0.155
Am-Ph	54.74	161.65	1.168	4.2	0.9	62	0.105
An-Au	-14.71	-140.30	0.653	1.6	0.8	33	0.004
An-Nu	-3.25	148.10	0.128	6.8	2.7	16	0.006
An-Nz	-35.12	90.90	0.453	4.0	1.0	-15	0.015
An-Pa	65.96	-85.36	0.857	0.6	0.4	88	0.011
An-Sa	84.64	-128.11	0.240	3.5	1.4	-80	0.011
An-So	-28.17	115.25	0.120	5.2	3.1	16	0.007
Ar-At	-26.08	-133.99	0.421	52.7	1.7	36	0.853
Ar-Eu	26.22	22.87	0.427	2.1	1.1	76	0.029
Ar-In	10.50	61.83	0.099	17.2	10.4	88	0.073
Ar-Nu	31.26	29.55	0.400	1.8	1.3	-85	0.030
Ar-So	21.06	28.62	0.441	1.8	1.0	55	0.029
At-Eu	26.62	34.37	0.833	20.0	0.7	-5	0.871
Au-Eu	12.57	46.19	0.640	1.3	0.5	-52	0.005
Au-In	-4.38	72.04	0.409	4.7	1.7	15	0.066
Au-Pa	61.40	6.16	1.080	0.6	0.4	62	0.008
Au-So	9.31	48.55	0.675	1.2	0.4	-57	0.004
Au-Su	8.64	58.16	0.743	6.5	0.9	-78	0.053
Ca-Na	75.45	-154.55	0.180	10.9	1.3	88	0.008
Ca-Nz	-48.76	73.13	0.382	5.6	3.1	21	0.020
Ca-Sa	52.83	-66.25	0.267	5.4	1.4	-5	0.021
Cs-Eu	41.90	-123.95	0.087	65.1	6.6	-50	0.006
Cs-Ph	56.57	167.80	1.123	4.2	0.9	66	0.063
Cs-Su	11.95	124.19	0.148	19.0	11.6	39	0.097
Eu-In	-28.56	-168.38	0.357	14.4	1.1	-89	0.033
Eu-Na	68.05	136.42	0.245	1.5	0.8	-38	0.004
Eu-Nu	18.23	159.99	0.062	9.5	3.7	-17	0.005
Eu-Ok	61.84	143.27	0.546	8.0	1.0	7	0.117
Eu-Pa	63.49	-78.34	0.904	0.6	0.4	78	0.006
Eu-Ph	55.79	162.02	1.063	2.6	0.5	45	0.040
Eu-Su	-8.76	104.55	0.181	5.4	2.2	-4	0.078
In-So	22.78	19.46	0.362	13.1	0.9	89	0.040
Na-Nu	-77.90	-75.23	0.213	2.0	1.2	64	0.004
Na-Ok	56.68	147.01	0.305	11.8	1.7	18	0.121
Na-Pa	50.38	-72.11	0.755	0.6	0.4	-79	0.004
Na-Sa	12.89	-50.43	0.171	2.7	1.3	-10	0.009
Na-Sr	-17.59	46.07	0.305	12.2	0.9	-27	0.103
Nu-Sa	62.67	-41.98	0.277	3.1	1.3	-2	0.006
Nu-So	-35.49	24.02	0.085	4.9	3.1	-19	0.005
Nz-Pa	55.41	-87.29	1.267	1.6	0.5	17	0.008
Nz-Sa	52.13	-91.18	0.633	3.1	1.2	8	0.009
Ok-Pa	27.81	-62.32	0.702	14.1	0.9	-19	0.047
Ok-Ph	47.58	175.21	0.538	10.5	1.7	79	0.120
Pa-Ph	4.60	138.27	0.874	1.7	0.6	-15	0.049
Pa-Sr	-45.18	85.40	0.950	3.5	0.4	-78	0.105
Pa-Su	-55.16	102.55	1.019	5.1	0.5	3	0.048
Ph-Su	-59.99	-1.24	1.047	6.8	0.8	-72	0.035

<sup>a</sup>The first plate rotates counterclockwise relative to the second plate around the listed rotation pole. Plate abbreviations are given in Table 1.

<sup>b</sup>Location of rotation pole,  $\lambda$  is latitude and  $\phi$  is longitude in decimal degrees.

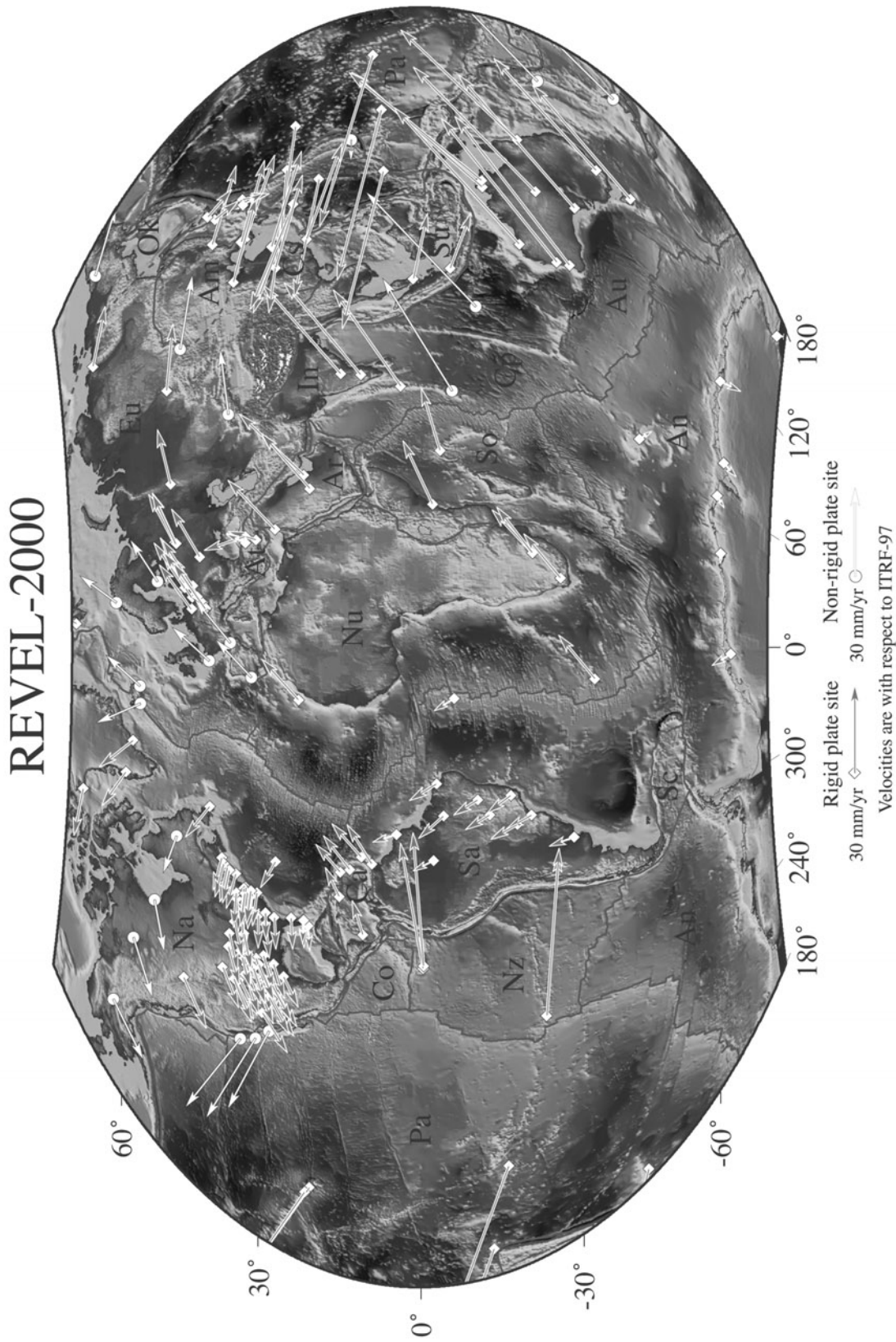
<sup>c</sup>Two-dimensional 1-sigma lengths in degrees of the semimajor  $\sigma_{\text{maj}}$  and semiminor  $\sigma_{\text{min}}$  axes of the pole error ellipse.

<sup>d</sup>The parameter  $\zeta$  is the azimuth of the semimajor ellipse axis in degrees clockwise from north.

misfit to the data for a given plate, as described by *Ward* [1990] and *Mao* [1998]. Relative angular velocities for pairs of adjacent plates are then derived by differencing the ITRF-97 angular velocities, with appropriate error propagation (Table 4). Tables 2, 3, and 4 constitute REVEL-2000.

[13] One advantage of our approach is that GPS site velocities can easily be compared to independent space geodetic data from SLR, VLBI, and DORIS, whose solutions are also available in ITRF-97. In addition, for plates with limited GPS data it is straightforward to incorporate data from these other techniques [e.g., *Norabuena et al.*, 1998, 1999]. (For these sites we use the ITRF-97 SINEX velocity and error solutions available from [\[lareg.ensg.ign.fr/ITRF\]\(http://lareg.ensg.ign.fr/ITRF\).\) We use these additional data for Arabia, Anatolia, and Nazca. One disadvantage of our approach is that three steps are required to derive a relative angular velocity \(namely, transform coordinate time series to ITRF-97 and derive site velocities; derive ITRF-97 plate angular velocities by inverting the site velocities; and difference these angular velocities to derive the relative angular velocities of adjacent plates\), whereas it is possible to go directly from loosely constrained \(pre-ITRF\) position estimates to a plate velocity-minimized frame, deriving a relative angular velocity in only two steps \[e.g., \*Kogan et al.\*, 2000\].](http://</a></p>
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[14] A plot of site velocities relative to ITRF-97 (Figure 4) superficially resembles the velocity field predicted by the



**Figure 4.** Winkel Tripel projection showing site velocities with respect to ITRF-97. Solid arrows (red) with diamonds are sites on rigid plate interiors; for clarity, sites on the Sierra Nevada block are omitted. Open arrows (white) with circles are selected sites in deforming plate interiors affected by GIA or other locations such that they are not used in the rigid plate definition. Plate boundaries are shown with blue lines, dashed where plate boundary is poorly defined. See Table 1 for plate abbreviations, except Sc, Scotia, and Cp, Capricorn. See color version of this figure at back of this issue.



**Table 5.** Comparison of Relative Angular Velocities for Selected Plate Pairs

Plate Pair <sup>a</sup>	$\phi$ , <sup>b</sup> °N	$\lambda$ , <sup>b</sup> °E	$\omega$ , deg/Myr	$\sigma_{\text{maj}}$ <sup>c</sup>	$\sigma_{\text{min}}$ <sup>c</sup>	$\zeta$ , <sup>d</sup> deg	$\sigma_{\omega}$ , deg/Myr	Author (Type of Data <sup>e</sup> )
Ar-Nu	31.26	29.55	0.400	1.8	1.3	-85	0.030	REVEL-2000
Ar-Nu	31.50	23.00	0.403	3.5	1.5	-74	0.045	<i>Chu and Gordon</i> [1998] (MM) <sup>f</sup>
Ar-Nu	32.59	23.70	0.400	—	—	—	—	<i>Jestin et al.</i> [1994] (MM, TA) <sup>g</sup>
Au-An	14.71	39.70	0.653	1.6	0.8	-33	0.004	REVEL-2000
Au-An	10.92	41.60	0.660	—	—	—	—	<i>Bouin and Vigny</i> [2000] (GP)
Au-An	11.89	42.82	0.650	3.1	1.0	-35	0.007	<i>Conder and Forsyth</i> [2000] (MM, TA) <sup>h</sup>
Au-An	9.80	43.20	0.650	4.4	2.6	20	0.01	<i>Larson et al.</i> [1997] (GPS) <sup>i</sup>
Au-An	13.20	38.20	0.650	1.3	1.0	-63	0.01	<i>DeMets et al.</i> [1994] (ES, MM, TA)
Ca-Na	75.45	-154.55	0.180	10.9	1.3	88	0.008	REVEL-2000
Ca-Na	69.10	-136.70	0.192	31.3	4.8	-7	0.036	<i>DeMets et al.</i> [2000] (GP)
Ca-Na	64.90	-109.50	0.214	14.6	1.5	-35	0.030	<i>DeMets et al.</i> [2000] (GP, MM, TA)
Ca-Na	68.40	-126.30	0.210	—	—	—	—	<i>Deng and Sykes</i> [1995] (ES)
Ca-Na	74.30	153.90	0.100	24.7	2.6	-52	0.030	<i>DeMets et al.</i> [1994] (ES, MM, TA)
Ca-Sa	52.83	-66.25	0.267	5.4	1.4	-5	0.021	REVEL-2000
Ca-Sa	61.90	-75.70	0.229	—	—	—	—	<i>Perez et al.</i> [2001] (GP)
Ca-Sa	51.50	-65.70	0.272	6.1	1.9	-9	0.023	<i>Weber et al.</i> [2001] (GP)
Ca-Sa	50.00	-65.30	0.180	14.9	4.3	-2	0.030	<i>DeMets et al.</i> [1994] (ES, MM, TA)
Cs-Eu	41.90	-123.95	0.087	65.1	6.6	-50	0.006	REVEL-2000
Cs-Eu	61.20	142.00	0.206	23.3	2.2	39	0.123	<i>Heki et al.</i> [1999] (GP)
Cs-Eu	64.84	156.74	0.120	—	—	—	—	<i>Holt et al.</i> [2000] (GP, QF)
Cs-Eu	-5.10	109.00	-0.190	6.2	0.7	-78	0.010	<i>Peltzer and Saucier</i> [1996] (ES)
Eu-Ok	61.84	143.27	0.546	8.0	1.0	7	0.117	REVEL-2000
Eu-Ok	53.02	142.09	0.405	13.5	3.4	-72	—	<i>Seno et al.</i> [1996] (ES)
Eu-Na	68.05	136.42	0.245	1.5	0.8	-38	0.004	REVEL-2000
Eu-Na	74.30	123.00	0.231	1.8	1.4	16	0.008	<i>Kogan et al.</i> [2000] (GP)
Eu-Na	78.50	122.00	0.230	8.2	4.9	-8	0.030	<i>Argus and Heflin</i> [1995] (GP)
Eu-Na	62.40	135.80	0.310	4.1	2.3	-11	0.010	<i>DeMets et al.</i> [1994] (ES, MM, TA)
In-Eu	28.56	11.62	0.357	14.4	1.1	89	0.033	REVEL-2000
In-Eu	25.60	11.10	0.440	—	—	—	0.030	<i>Paul et al.</i> [2001] (GP)
In-Eu	29.78	7.51	0.353	—	—	—	—	<i>Holt et al.</i> [2000] (GP)
In-Eu	24.40	17.70	0.510	8.8	1.8	-79	0.050	<i>DeMets et al.</i> [1994] (ES, MM, TA)
Na-Pa	50.38	-72.11	0.755	0.6	0.4	-79	0.004	REVEL-2000
Na-Pa	51.50	-73.70	0.765	2.0	1.0	-85	0.016	<i>DeMets and Dixon</i> [1999] (GS)
Na-Pa	50.00	-76.10	0.777	0.8	0.6	65	0.007	<i>DeMets and Dixon</i> [1999] (MM, TA)
Na-Pa	49.60	-84.30	0.830	2.0	1.0	-86	0.020	<i>Larson et al.</i> [1997] (GP) <sup>i</sup>
Na-Pa	49.10	-73.00	0.790	4.1	2.2	-83	0.030	<i>Argus and Heflin</i> [1995] (GP)
Na-Pa	48.70	-78.20	0.750	1.3	1.2	-61	0.010	<i>DeMets et al.</i> [1994] (ES, MM, TA)
Na-Sa	12.89	-50.43	0.171	2.7	1.3	-10	0.009	REVEL-2000
Na-Sa	12.00	-50.60	0.130	6.8	3.0	-19	0.03	<i>Dixon and Mao</i> [1997] (GP)
Na-Sa	11.10	-53.30	0.290	76.6	3.9	37	0.08	<i>Larson et al.</i> [1997] (GP) <sup>i</sup>
Na-Sa	6.50	-55.60	0.280	8.3	7.4	-55	0.12	<i>Argus and Heflin</i> [1995] (GP)
Na-Sa	16.30	-58.10	0.150	5.9	3.7	-9	0.01	<i>DeMets et al.</i> [1994] (ES, MM, TA)
Nu-So	-35.49	24.02	0.085	4.9	3.1	-19	0.005	REVEL-2000
Nu-So	-27.30	36.20	0.089	6.9	5.5	38	0.004	<i>Chu and Gordon</i> [1999] (MM, TA) <sup>f</sup>
Nu-So	-55.73	19.76	0.052	—	—	—	—	<i>Jestin et al.</i> [1994] (TA, MM) <sup>g</sup>
Ph-Pa	-4.60	-41.73	0.874	1.7	0.6	15	0.049	REVEL-2000
Ph-Pa	-7.51	-42.81	1.009	—	—	—	—	<i>Wei and Seno</i> [1998] (ES)
Ph-Pa	-1.24	-45.81	1.000	—	—	—	—	<i>Seno et al.</i> [1993] (ES)

<sup>a</sup> The first plate rotates counterclockwise relative to the second plate around the listed rotation pole. Plate abbreviations are given in Table 1.

<sup>b</sup> Location of rotation pole,  $\lambda$  is latitude and  $\phi$  is longitude in decimal degrees.

<sup>c</sup> Two-dimensional 1-sigma lengths in degrees of the semimajor  $\sigma_{\text{maj}}$  and semiminor  $\sigma_{\text{min}}$  axes of the pole error ellipse.

<sup>d</sup> The parameter  $\zeta$  is the azimuth of the semimajor ellipse axis in degrees clockwise from north.

<sup>e</sup> ES, earthquake slip vectors; GP, GPS; MM, marine magnetic; QF, Quaternary fault slip rates; TA, transform azimuth; VL, VLBI.

<sup>f</sup> Estimated from their covariance matrix.

<sup>g</sup> The best enforced closure solution and velocity reduced by 4.4% to match Hilgen timescale.

<sup>h</sup> Based on 1 Ma “Jaramillo event.”

<sup>i</sup> Originally published as reverse plate pair.

NUVEL-1 No-Net-Rotation (NUVEL-1-NNR) velocity field [*Argus and Gordon*, 1991] or its updated version, NUVEL-1A-NNR [*DeMets et al.*, 1994], because of the way ITRF is defined. Inasmuch as the NUVEL-1A geologic model may be a biased representation of Recent plate motions for some plate pairs, as described earlier, it may be inappropriate to use the NNR reference frame based on NUVEL-1A for geodetic observations. Recent versions of the ITRF are believed to be free of biases associated with a particular plate motion model [*Sillard et al.*, 1998]. However, to be conservative, in our approach, ITRF-97 is treated as an arbitrary initial reference frame, differenced out in

the subsequent calculation of relative angular velocities. Hence any bias in the ITRF should not influence our results. The success of our approach can be assessed in two ways:

1. The orientation of residual velocities on a given plate should be random, and their magnitude should be similar to the site velocity uncertainty.

2. Independent geologic and geodetic estimates along a given plate boundary should agree with the predictions of our angular velocities.

[15] Both tests are applied in this paper and suggest that for the most part our approach is valid. Briefly, residual velocities for most



**Figure 5.** Lambert azimuthal equal-area projection showing all space geodetic site locations used in this study for the area around the Indian Ocean. Earthquake epicenters (magnitude  $>4.5$ , depth  $<50$  km, between 1973 and 2000) suggest plate boundaries. Sites labeled with diamonds and upper case letters are assumed to lie on the rigid plate interior and are used to define the plate's angular velocity (Table 2). Sites labeled with circles and lowercase letters are not used in the rigid plate definition (Table 2). See Table 1 for plate abbreviations, except Cp, Capricorn. For clarity, site names in some areas have been omitted, e.g., Europe.

plates are randomly oriented in azimuth and have magnitudes that are about the level expected given observational error. Exceptions are explicitly discussed below.

[16] As our new velocity data incorporate a longer time span, additional sites, and improved analytical techniques, the angular velocities presented here are more precise than our previous published estimates for individual plate pairs. However, the major conclusions of those studies remain unchanged [Dixon and Mao, 1997; Dixon *et al.*, 1998; DeMets and Dixon, 1999; Norabuena *et al.*, 1999; Dixon *et al.*, 2000a, 2000b; Weber *et al.*, 2001].

### 3. Results

[17] Our results are summarized in Tables 1, 2, 3, and 4 and Figures 4–21. Site locations are displayed in the global map figures (Figures 5, 13, and 17), along with seismicity (magnitude  $>4.5$  and depth  $<50$  km) for the period 1973–2000, from the National Earthquake Information Center (<http://neic.usgs.gov>), which generally outlines the appropriate plate boundary. Table 1 lists the individual site velocities in ITRF-97 grouped by plate, as well as site positions, total time span of data, number of data (number of daily position estimates), weighted root mean square scatter (WRMS) of the data about a best fit line (an indicator of data quality), and the residual velocities

predicted from the best fitting ITRF-97 angular velocity for each plate (Table 2).

[18] Table 2 lists the best fitting angular velocity for each plate or block with respect to ITRF-97, the mean rate and RMS of the velocity residuals, and the  $\chi^2$  per degree of freedom. Table 3 lists the corresponding Cartesian coordinates and covariance matrix. Table 4 lists the relative angular velocities for pairs of plates sharing a common boundary. Table 5 compares a subset of these angular velocities to some previously published values.

## 4. Discussion

### 4.1. Plate Rigidity and Error Model

[19] Rigid plate behavior on timescales of several million years is demonstrated by the success of rigid plate geologic models, and the exceptions are well described and spatially limited [Gordon, 1998]. Space geodesy can test the assumption of rigid plates on shorter timescales [Argus and Gordon, 1996; Dixon *et al.*, 1996]. The existence of intraplate seismicity such as the New Madrid seismic zone [Nuttli, 1973; Schweig and Ellis, 1994; Weber *et al.*, 1998; Newman *et al.*, 1999] argues that some local intraplate deformation must occur. Plate rigidity over several years at the level of a few millimeters per year is suggested by the generally good agreement between geologic models and space geodetic

estimates of plate motion [Smith *et al.*, 1990]. On the other hand, the spatial sampling afforded by techniques such as VLBI and SLR is limited, and until recently, the precision of GPS, which in principle can provide better spatial sampling, has also been limited. Dixon *et al.* [1996] noted that with the data then available, GPS data lacked the sensitivity to detect a clear postglacial rebound signal and provided only a crude upper limit to plate rigidity, as apparent deviations from a rigid plate model (e.g., the velocity residuals in Tables 1 and 2) mainly reflect GPS velocity uncertainties rather than true nonrigid plate processes. While the current data set is a significant improvement relative to the data available in 1996, and apparently has some sensitivity to postglacial rebound (see section 4.2), the same limitation still applies, although we can now place a tighter upper bound on plate rigidity from the GPS data. Useful constraints are available in the current data set for the North American, Eurasian, and possibly Australian plates. For the remaining plates, the number of data is too small to derive rigorous estimates of plate rigidity, although we note that the velocity data fit a rigid plate model at about the expected level for most of the plates studied (Figure 2b).

[20] Using all 84 available sites to define stable North America, comprising sites on the plate interior that have observations spanning as short as 2.0 years, the mean rate residual is 0.99 mm/yr. Below, we describe a 64-site solution that eliminates sites most sensitive to postglacial rebound, resulting in a mean rate residual of 0.86 mm/yr. In either case the residual magnitude is roughly the same as the GPS velocity error, suggesting that the rigid plate model is appropriate and that the residuals likely reflect GPS observation errors rather than nonrigid plate processes, as found in earlier studies [e.g., Dixon *et al.*, 1996]. Since GPS velocity errors are a strong function of total observing time, one way to check this conclusion is to restrict the data set to sites with longer time spans and to investigate the effect on residuals. The residual magnitudes as a function of mean observing time are consistent with the error model illustrated in Figure 3. For example, if we base the angular velocity estimate for North America on the 36 sites in our database that are not sensitive to postglacial rebound and have time spans of 4 years or longer, the mean rate residual is 0.71 mm/yr, compared to 0.86 mm/yr for the larger data set, supporting the inference that the rate residuals reflect GPS velocity errors rather than nonrigid plate processes.

[21] Rate residuals are positively biased (*Argus and Gordon* [1996] and *Argus et al.* [1999] give a complete discussion) implying that the actual level of plate nonrigidity could be higher than the mean rate residual. If we arbitrarily assume that an upper bound on plate rigidity is 3 times the mean rate residual, then the GPS data set composed of 4 years or longer time series suggests that the stable interior of North America is rigid to better than 2.2 mm/yr. Presumably, this bound will be better constrained in future years as GPS velocity uncertainties decline further and new sites are added. Of the 19 plates and continental blocks examined here, all but three (Anatolia, Arabia, Nazca) have mean rate residuals of 1.8 mm/yr or less (Table 2). These three plates or blocks have a very limited distribution of sites, suggesting that their angular velocities are not well determined. The main point for this paper is that the velocity residuals, whether reflecting true nonrigid plate processes, local site effects, or (our preferred explanation) GPS velocity uncertainties, are small enough for the rigid plate approximation to be assumed valid for the data set under consideration.

[22] The relatively uniform nature of space geodetic data makes it feasible to apply a consistent error model, facilitating statistical tests of plate rigidity. If our error model for GPS velocities is approximately correct,  $\chi^2$  per degree of freedom (hereafter  $\chi_v^2$ ) for the individual plates should approximately equal 1.0, provided that the data fit the rigid plate model, and that the data set is large enough to be statistically representative. The latter criterion is met for the North American plate (128 data;  $\chi_v^2 = 1.05$ ) and the

Eurasian plate (30 data;  $\chi_v^2 = 1.02$ ). However, the majority of plates tested also appear to approximately satisfy this “rule of thumb,” despite the relatively sparse data (Figure 2b). All but three plates or blocks (Amuria, Australia, Okhotsk) have  $\chi_v^2 < 2.0$ , and two of these (Amuria and Okhotsk) are probably related to limited data (Table 2). Since the error model is derived independently of any rigid plate criteria, this tends to confirm the joint hypothesis that the great majority of plates and blocks tested are rigid within the velocity uncertainty, that the sites used to define the various plates lie on the rigid portion of the plate interior, and that the error model is approximately correct. The fact that  $\chi_v^2$  for North America and Eurasia is slightly larger than 1.00 might reflect the fact that we have neglected random walk noise. However, the effect is small and is ignored here.

[23] For several plates the number of space geodetic data is small enough that the solutions are sensitive to outliers at one or two sites, the statistical effect of which is diminished when a large number of sites is available. For example, Amuria and Okhotsk, with only three sites and five sites, respectively, have  $\chi_v^2 = 2.3$  and 4.4, suggesting that these data are not well fit by the rigid plate model. Possible explanations include that (1) the number of data on the plate are too small to be statistically representative; (2) one or more site velocities have a systematic error not reflected in the error model; (3) our choice of sites is inappropriate, and one or more sites lie in a deforming boundary zone rather than the rigid plate or block interior; and (4) the plate is not rigid. We suspect that explanations 1, 2, and 3 together explain such misfits, reflecting limitations in available data.

#### 4.2. Glacial Isostatic Adjustment

[24] The lithosphere in North America, Eurasia, and Antarctica is isostatically adjusting from mass loading and unloading of the last glaciation, reflecting the delayed response associated with viscous flow in the Earth’s mantle. This adjustment imparts both vertical and horizontal motions to the surface velocity field [e.g., *Peltier*, 1998a] and needs to be considered when using geodetic data to derive plate motion models because the velocity effects are not representative of longer-term motions. That part of the surface velocity field due to glacial isostatic adjustment (GIA) is a function of both ice loading history and mantle viscosity structure, neither of which is well known, so the corresponding model predictions can vary significantly [e.g., *Argus et al.*, 1999; *Mitrovica et al.*, 2000]. However, most models agree on the region of maximum glacial isostatic effect, even if they do not agree on the magnitude or direction of predicted motion. Rather than correcting each site velocity by a specific model prediction, we take a more conservative approach and simply eliminate the subset of sites most likely to be affected by GIA. For North America, maximum GIA-related motion is centered in a broad region around Hudson Bay. We use the ICE-4G model [*Peltier*, 1994] to define sites with horizontal rates due to GIA  $\geq 0.7$  mm/yr and compare solutions with and without this data subset. Thus we compare a solution for North America using all available data (84 sites,  $\chi_v^2 = 1.63$ ) versus a model that omits 20 sites most affected by GIA (algo, chb1, chur, det1, dubo, flin, kew1, mil1, neb3, nrc1, oro\_, sag1, sch2, stb1, stp1, vcap, whp1, wis1, yell, you1) giving a 64-site model with  $\chi_v^2 = 1.05$ . Of these 20 sites, 9 exhibit statistically significant uplift (greater than one standard error), and 5 exhibit uplift rates greater than 3.0 mm/yr (Table 1). Sites may also subside due to GIA, associated with collapse of the peripheral bulge, but associated horizontal velocities are thought to be small compared to our observational error.

[25] For Eurasia we compared the solution using all available sites (19 sites,  $\chi_v^2 = 3.73$ ) versus a solution that omits four sites on the Fennoscandian platform (kiru, mets, onsa, trom) most likely to be affected by GIA. The omitted sites were chosen using the same criteria for horizontal motions as for North America



[Peltier, 1998b], approximately equivalent to eliminating sites with vertical motion  $>3$  mm/yr, as predicted by the uplift model of Davis *et al.* [1999]. In this case the fit is significantly improved (15 sites,  $\chi_v^2 = 1.02$ ). We conclude that the current GPS velocity field is sensitive to GIA and that correcting for this effect or eliminating sites sensitive to it is important for obtaining accurate plate angular velocity estimates for comparison to geologic values. For Antarctica we cannot do this as all sites are experiencing GIA uplift (see Table 1).

#### 4.3. Comparison to Geologic and Geodetic Data

[26] In this section we discuss the criteria for inclusion or omission of key sites by plate (listed alphabetically, except South America, included in the Nazca plate discussion), compare our results to independent data and the NUVEL-1A geologic model [DeMets *et al.*, 1994], and discuss some implications. The discussion below focuses on those plates or blocks where limited available data make the results sensitive to the selection criteria, where differences appear to exist with NUVEL-1A, and on the relative velocities of several of the “new” plates or continental blocks that have not been included in previous global models. We omit discussion of plates or blocks where our data are limited or where our results do not differ significantly from previous publications: Anatolia [Reilinger *et al.*, 1997; McClusky *et al.*, 2000], Okhotsk [Seno *et al.*, 1996; Wei and Seno, 1998], Sierra Nevada [Dixon *et al.*, 2000a], and Sunda [Walpersdorf *et al.*, 1998; Rangin *et al.*, 1999; Chamot-Rooke and Le Pichon, 1999].

**4.3.1. Amuria.** [27] Angular velocities for the Amurian plate have been reported by Zonenshain and Savostin [1981], Wei and Seno [1998], Heki *et al.* [1999], and Holt *et al.* [2000]. Following Heki *et al.* [1999], we use three sites to define this plate, DAE\*, SUWN, and VLAD (Figure 5). The limited geographic distribution results in correspondingly large uncertainties. Site bjfs is not included in our rigid plate definition since it is located in an area of relatively high seismicity and active faulting [Shen *et al.*, 2000] (the 4-site solution including bjfs,  $62.64^\circ\text{N}$ ,  $-128.76^\circ\text{E}$ ,  $0.319^\circ/\text{Myr}$ ,  $\sigma_{\text{maj}} = 17.1$ ,  $\sigma_{\text{min}} = 1.3$ ,  $\sigma_{\omega} = 0.036^\circ/\text{Myr}$ , mean rate residual (MMR) = 1.2 mm/yr, is similar to the 3-site solution listed in Table 2 and improves  $\chi_v^2$  from 2.30 to 1.46).

[28] The formation of the Baikal rift reflects relative motion between Amuria and Eurasia [Zonenshain and Savostin, 1981]. Our relative angular velocity predicts south-southeast extension across the rift at  $6-7$  mm/yr. At  $47.5^\circ\text{N}$ ,  $106.5^\circ\text{E}$ , close to where Calais and Amarjargal [2000] measure a velocity of  $6.4 \pm 1.6$  mm/yr at an azimuth of  $125 \pm 30^\circ$  from continuous GPS data, we predict (Amuria relative to Eurasia)  $7.0 \pm 3.2$  mm/yr,  $165 \pm 19^\circ$ .

**4.3.2. Antarctica.** [29] We define the Antarctic plate using seven sites (Table 1), excluding ohig and palm because of their proximity to the complex plate boundary with the Scotia plate [Pelayo and Wiens, 1989; Klepeis and Lawver, 1996]. The Antarctic-Australia and Antarctic-Pacific boundaries are spreading ridges, and their velocities are well determined in geologic models such as NUVEL-1A. Pacific and Australia velocities are also well determined in REVEL-2000 (Tables 1, 2, and 3). Any difference in velocity between the two models probably indicates real velocity changes. Figures 6 and 7 indicate that the rates and azimuths predicted by the two models for both plate pairs are virtually identical along most of the respective plate boundaries. For Australia-Antarctica, REVEL-2000 and NUVEL-1A agree to better than 0.6 mm/yr in rate for  $>80\%$  of the entire plate boundary, implying a remarkable steadiness of plate motion for this plate pair over the last 3 Myr.

**4.3.3. Arabia.** [30] We define the Arabian plate using a combination of two GPS sites (BAHR, KATZ), and one SLR site (7832) (Figure 5). KATZ is within 30 km of the active Dead Sea fault and thus does not meet our criterion for minimum distance from the plate boundary. However, the amount of strain

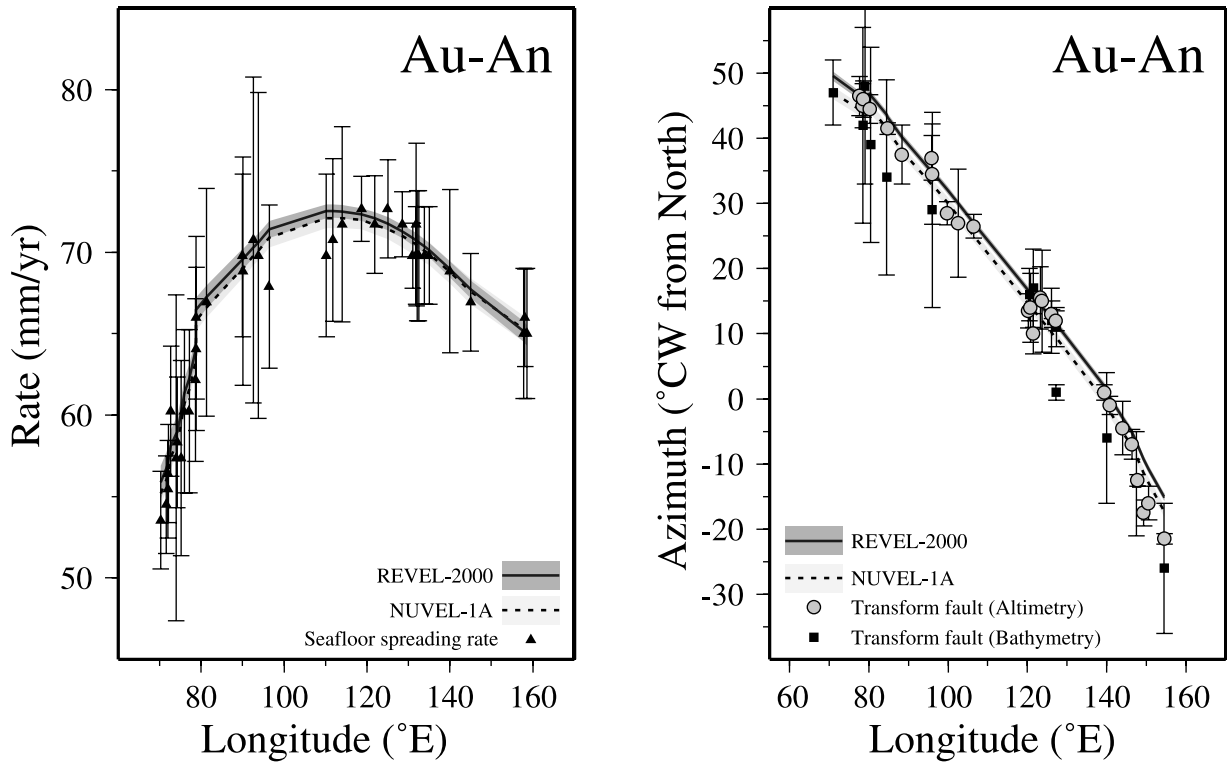
accumulation is believed to be small here [Pe’eri *et al.*, 2002]. The north striking Dead Sea fault is a left-lateral strike-slip fault separating Arabia and the Sinai block, a small continental block that may move relative to Nubia. Young ( $<140$  kyr) offset geomorphic features along the Dead Sea fault at  $30.8^\circ\text{N}$ ,  $35.4^\circ\text{E}$  indicate a slip rate of  $4 \pm 2$  mm/yr at a strike of  $028^\circ$  [Klinger *et al.*, 2000]. Our predicted Arabia-Nubia velocity at this location is very similar,  $3.9 \pm 0.8$  mm/yr at an azimuth of  $009 \pm 10^\circ$ , implying very slow motion of Sinai relative to Nubia.

[31] If motion of Arabia relative to Nubia has been steady over the last few million years, we expect agreement between the REVEL-2000 prediction and the rate of seafloor spreading across the Red Sea. However, measured spreading rates averaged over the last 3 Myr [Chu and Gordon, 1998], as well as the model rates of Jestin *et al.* [1994] based on a similar time period, are both systematically higher than REVEL-2000 (Figure 8 and Table 5). Similarly, NUVEL-1A’s predicted Arabia-Eurasia rate is higher than REVEL-2000 (Figure 9). McClusky *et al.* [2000] obtained a similar result. These rate differences may reflect gradual slowing of the Arabian plate as it moves north and collides with Eurasia. Associated crustal thickening forms the Zagros [Alavi, 1994] and Caucasus Mountains, increasing gravitational body forces that oppose Arabia’s northward motion, perhaps resulting in gradual slowing of Arabian-Eurasia convergence and Red Sea spreading.

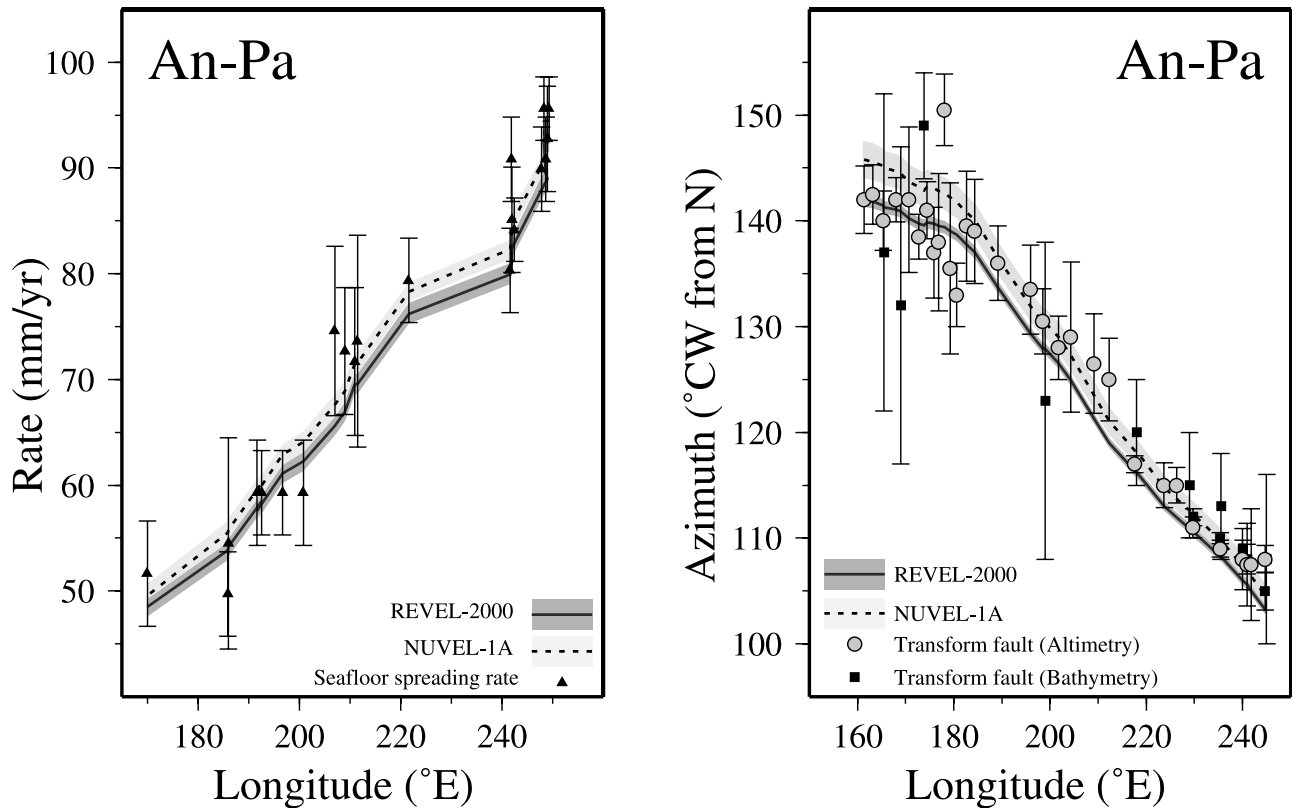
**4.3.4. Australia.** [32] We define Australian plate motion using 11 sites on Australia and Tasmania (Table 1 and Figure 10). Sites noum and auck have velocities relative to stable Australia that are larger than the corresponding velocities of stable interior sites, perhaps reflecting strain accumulation associated with nearby subduction zones, suggesting that they should not be used for the Australian plate definition. The parameter  $\chi_v^2$  is 3.18 for a 13-site solution that includes noum and auck versus 2.78 for the 11-site solution;  $\chi_v^2$  is 3.31 for a 12-site solution that includes auck. However, even omitting noum and auck results in misfits and a  $\chi_v^2$  that is higher than expected. The velocity residual pattern is very similar to the direction of maximum horizontal compression in the region. Recent compilations of stress data [Hillis and Reynolds, 2000] suggest east-west compression in western Australia, rotating to northwest-southeast compression in southeastern Australia and northeast-southwest compression in northeast Australia (e.g., in vicinity of TOW2) (Figure 10). This may explain why  $\chi_v^2$  for this plate is high. In other words, the rigid plate assumption may not be strictly valid for the continental portion of the plate, and Australia may be undergoing intraplate deformation in response to these compressional stresses. Alternatively, we may have underestimated the site velocity errors for Australia. We prefer the former explanation because the error model gives consistent results for most other plates (Figure 2b).

[33] Separate Indian and Australian plates have been recognized for some time [Wiens *et al.*, 1985; DeMets *et al.*, 1988], and a Capricorn plate has also been proposed in the central western Indian Ocean with a broad, diffuse boundary [Royer and Gordon, 1997; Gordon *et al.*, 1998]. We are unable to calculate relative angular velocities for plate pairs involving the Capricorn plate because of insufficient geodetic data. However, we can use the  $\chi^2$  test to evaluate the fit of the velocity data for coco (Cocos Island, on the southern edge of India-Australia boundary), dgar (Diego Garcia, near the inferred boundary between the India and Capricorn plates), HYDE (Hyderabad, India), IISC (Bangalore, India), and MALD (Male, Maldives) to our rigid Australian plate model, with 11 sites and  $\chi_v^2 = 2.78$  (Figure 5). The  $\chi_v^2$  for the 11 + 5 sites is 15.25, i.e., significantly higher, indicating a very poor fit. If we exclude HYDE, IISC, and MALD (11 + 2 sites),  $\chi_v^2$  drops to 2.74, consistent with separate Indian and Australian plates (e.g., using the  $F$  ratio test of Stein and Gordon [1984]). However, when we exclude either dgar or coco in an 11 + 1 site solution for Australia, we get little or no improvement (excluding dgar gives  $\chi_v^2 = 2.72$ , excluding coco gives  $\chi_v^2 = 2.86$ ). The small change from excluding

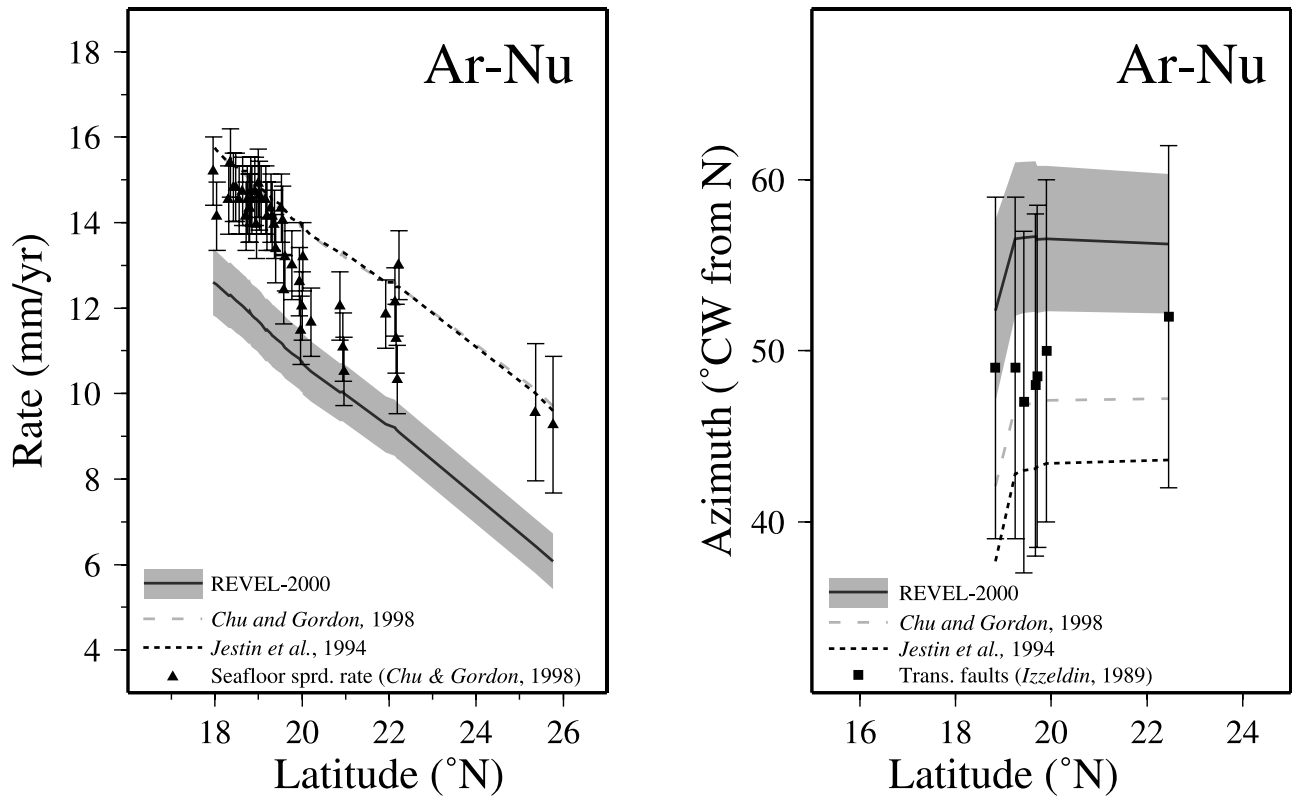




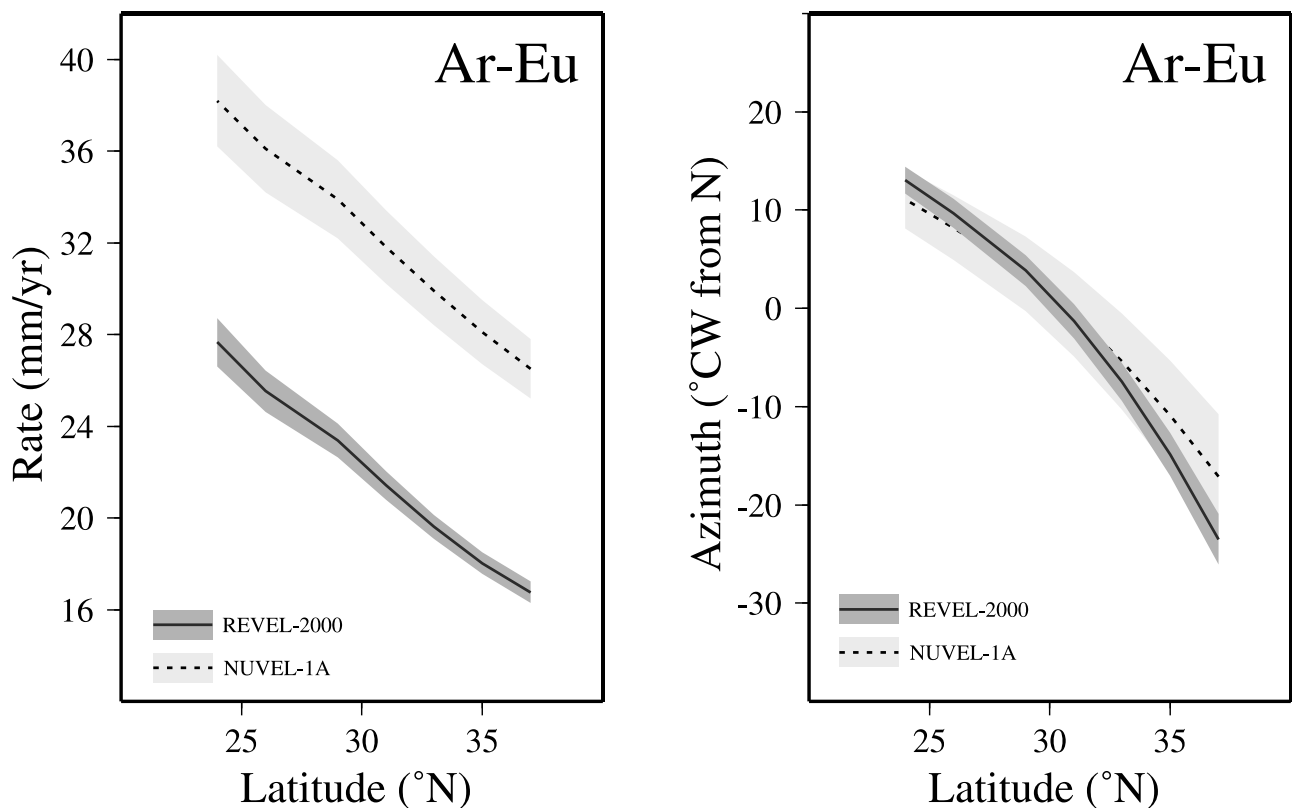
**Figure 6.** Measured seafloor spreading rates and transform fault azimuths for Australia-Antarctica compared to predicted rate and azimuth at the same location from REVEL-2000 (this study) and NUVEL-1A [DeMets *et al.*, 1994]. Spreading rates are from DeMets *et al.* [1990] corrected according to DeMets *et al.* [1994]. Transform azimuths from bathymetric measurements are from DeMets *et al.* [1990] unless otherwise noted. Transform azimuths from altimetry are from Spitzak and DeMets [1996]. Dark shading indicates  $\pm 1$  standard error for REVEL-2000. Light shading indicates  $\pm 1$  standard error for NUVEL-1A.



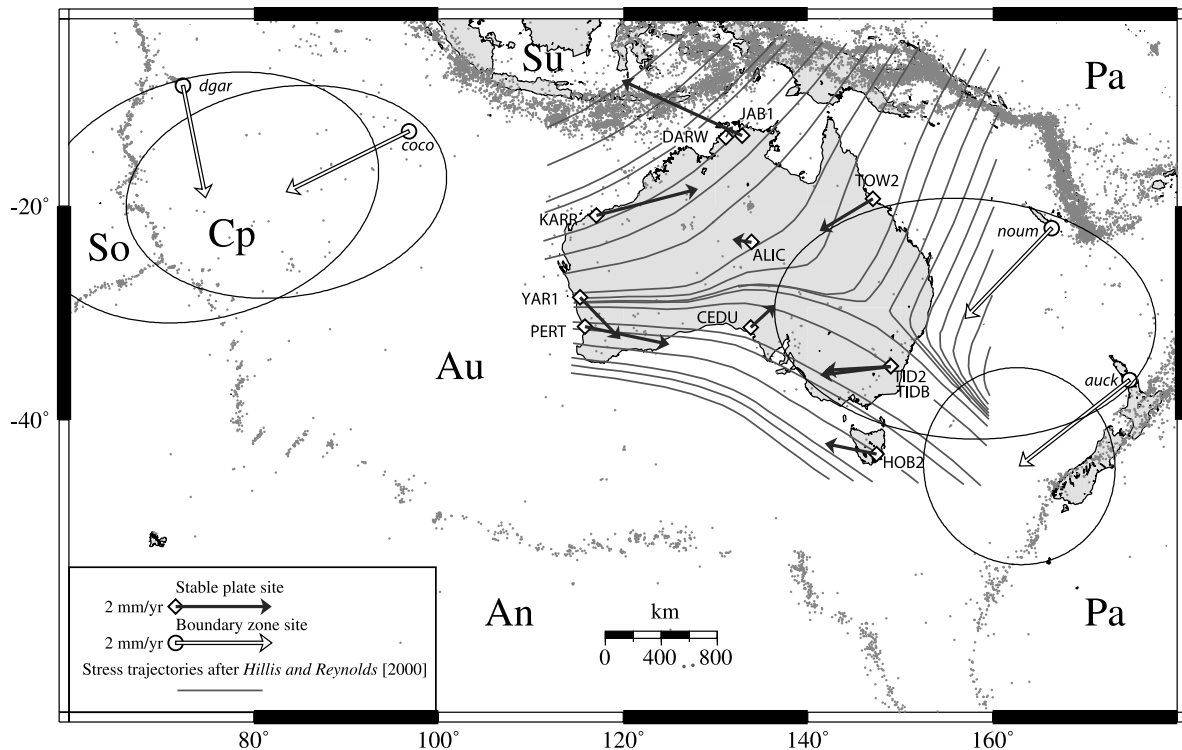
**Figure 7.** Similar to Figure 6, for Antarctica-Pacific.



**Figure 8.** Similar to Figure 6, for Arabia-Nubia. *Jestin et al.* [1994] pole is their best enforced closure pole (their Table 4) and has been converted to Hilgen timescale by reducing the angular velocity by 4.4%. Transform fault azimuth data of *Izzeldin* [1989] are only those transform faults that offset at least two marine magnetic anomalies (his Table 1).



**Figure 9.** Similar to Figure 6, for Arabia-Eurasia. Site predictions are calculated along a line that approximates the plate boundary, subparallel to the north shore of the Persian Gulf (37°N, 44°E to 24°N, 62°E).



**Figure 10.** Lambert azimuthal equal-area projection showing space geodetic site locations for the Australian plate and earthquake epicenters (magnitude  $>4.5$ , depth  $<50$  km, between 1973 and 2000), indicating plate boundaries. Sites with diamonds and upper case letters are assumed to lie on the stable plate interior and are used to define the plate's angular velocity. Their residual velocities with respect to the plate interior are shown by solid arrows (Table 2), and for clarity, their error ellipses are omitted. Sites with circles and lowercase letters lie in deforming boundary zones and are not used in the rigid plate definition. Their residual velocities with respect to the plate interior are shown with open arrows (Table 2) and error ellipses representing two-dimensional 95% confidence regions. Stress trajectories (orientation of maximum horizontal compression) from Hillis and Reynolds [2000] are superimposed, showing a similar pattern to residual velocities in the plate interior. See Table 1 for plate abbreviations, except Cp, Capricorn.

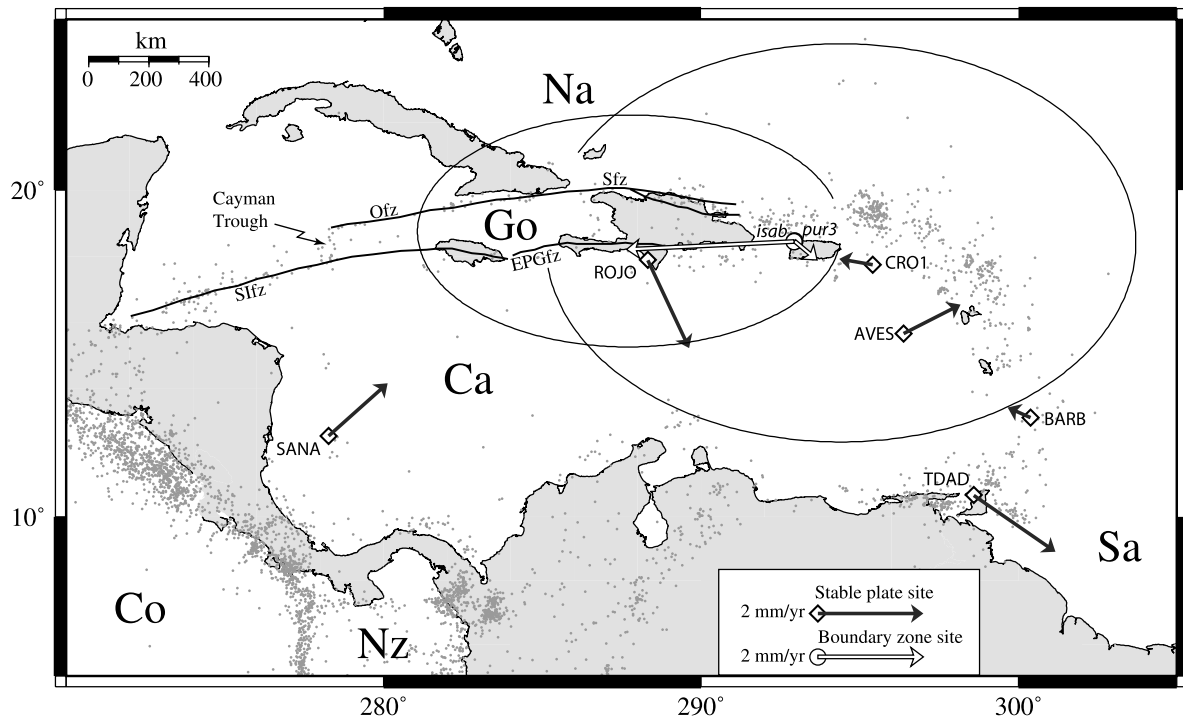
coco or dgar reflects the fact that these two sites have velocities that are negligibly different from stable Australia, consistent with very slow relative motion between Capricorn and Australia [Royer and Gordon, 1997; Conder and Forsyth, 2001].

**4.3.5. Caribbean.** [34] The Caribbean plate has been the focus of kinematic studies for at least 25 years [Jordan, 1975; Stein *et al.*, 1988; Deng and Sykes, 1995; DeMets *et al.*, 2000]. Geologic estimates of the velocity of the Caribbean plate with respect to its neighbors are hampered by the paucity of relevant data (especially rate data from spreading centers) and the tectonic complexity of the boundary zones around the plate (see discussions by DeMets *et al.* [1990] and Dixon *et al.* [1991b]). The velocity data set used here is similar to that presented by Weber *et al.* [2001], adding only additional position data at the continuous sites BARB and CRO1. Also, we take a slightly more conservative approach, omitting two sites in Puerto Rico (isab, pur3) because of the possibility of independent motion of the Puerto Rico block [Jansma *et al.*, 2000]. The resulting Caribbean-South America angular velocity is nevertheless essentially identical to that presented by Weber *et al.* [2001] (Table 5). The Caribbean-North America angular velocity is similar to that presented by DeMets *et al.* [2000], although the current vector is constrained by additional position data at SANA, ROJO, and CRO1, which improves their velocity estimate, as well as by velocities from new sites at BARB and TDAD. As pointed out by Dixon *et al.* [1998], Pollitz and Dixon [1998], DeMets *et al.* [2000], Weber *et al.* [2001], and Perez *et al.* [2001], the motion of the Caribbean plate with respect to both North and South America is considerably

faster than predicted by NUVEL-1A, probably reflecting systematic errors in the geologic model (Table 5).

[35] With the benefit of hindsight, it is clear why NUVEL-1A systematically underestimates the speed of the Caribbean plate (the predicted azimuths for NUVEL-1A and REVEL-2000 are very similar along much of the Caribbean boundary). Sykes *et al.* [1982], Rosencrantz and Mann [1991], and Mann *et al.* [1995] discuss the importance of the Gonave microplate, separating the North American and Caribbean plates. This microplate is defined by the Cayman spreading center on the west [Macdonald and Holcombe, 1978; Rosencrantz *et al.*, 1988], the Oriente-Septentrional fault zone on the north, and the Enriquillo-Plantain Garden fault zone on the south [Leroy *et al.*, 2000; Pubellier *et al.*, 2000] (Figure 11). Dixon *et al.* [1998] estimated  $8 \pm 4$  mm/yr of motion along the Enriquillo fault zone in the Dominican Republic, a significant fraction of overall North America-Caribbean motion. NUVEL-1A underestimates Caribbean-North America (and, by implication, Caribbean-South America) motion by roughly the slip rate on this fault because this rate represents that portion of plate motion not accommodated on the Cayman spreading center, NUVEL-1A's only Caribbean rate datum. REVEL-2000 provides a more accurate estimate of Caribbean-North America and Caribbean-South America motion, not only for the decade timescale but perhaps for the last few million years as well, because it represents the total relative plate motion, not just motion accommodated at the spreading center.

**4.3.6. Eurasia.** [36] We define the Eurasian plate using 15 sites (Figure 5 and Table 1) and obtain a well-defined



**Figure 11.** Similar to Figure 10, for the Caribbean plate. EPGfz, Enriquillo-Plantain Garden fault zone; Go, Gonave block; Ofz, Oriente fault zone; Sfz, Septentrional fault zone; Sifz, Swan Island fault zone.

solution ( $\chi_v^2 = 1.02$ , mean rate residual of 0.71 mm/yr). We exclude all sites south of significant seismicity reflecting the Nubia-Eurasia collision (e.g., south of the Pyrenees we exclude ebre, and vill [Ribeiro *et al.*, 1996; Herraiz *et al.*, 2000]), all sites in or south of the Alps and south of the Carpathians (geno, noto, penc, sofi) [Calais, 1999; Grenerczy *et al.*, 2000], all sites in central Asia south of 45°N because they may be affected by the India-Eurasia collision (kit3, pol2, sumk, urum), and all sites east of 130°W that may be on the North American plate or in the deforming zone between North America and Eurasia (bili). We also exclude sites in or near the Rhine Graben and Roer Graben (dour, kosg, wsrt) and sites west of the Rhine Graben and south of the Roer Graben (brst, brus, hers, mans, mlvl, sjdv, toul) similar to Nocquet *et al.* [2001]. These grabens were active in late Pliocene time and are also the current locus of significant seismicity [Plenefisch and Bonjer, 1997]. In addition, irkt is excluded because it is <100 km from the active Baikal Rift [Doser, 1991; Delvaux *et al.*, 1997] (see section 4.3.1), and yakz is excluded because of problems with its time series that we do not understand (our time series for this site shows a large seasonal variation in all three components). We also omitted kiru, mets, onsa, and trom due to sensitivity to GIA, as discussed earlier. Site brst shows a large residual motion with respect to Eurasia, possibly reflecting neotectonic effects in the area [van Vliet-Lanoe *et al.*, 1997]. A Eurasia solution that excludes TIXI, near the possible plate boundary with North America, results in an angular velocity that is negligibly different (14-site solution 58.37°N, -102.06°E, 0.258°/Myr,  $\sigma_{\text{maj}} = 1.7$ ,  $\sigma_{\text{min}} = 0.4$ ,  $\sigma_{\omega} = 0.004$ °/Myr,  $\chi_v^2 = 1.10$ , MRR = 0.7 mm/yr).

[37] Inclusion of the available six sites in stable western Europe (west of the Rhine Graben) in the Eurasia plate solution increases misfit ( $\chi_v^2 = 1.70$  for 21 sites compared to our preferred 15-site solution with  $\chi_v^2 = 1.02$ ), but the difference is small enough that separate western Europe and Eurasian plates are not supported by an  $F$  ratio test [Stein and Gordon, 1984] (angular velocity for the six sites west of the Rhine Graben with respect to ITRF-97 is 45.64°N, -117.08°E, 0.215°/Myr,  $\sigma_{\text{maj}} = 18.1$ ,  $\sigma_{\text{min}} = 0.9$ ,  $\sigma_{\omega} = 0.016$ °/Myr,  $\chi_v^2 = 1.46$ , MRR = 0.8 mm/yr). The two plates

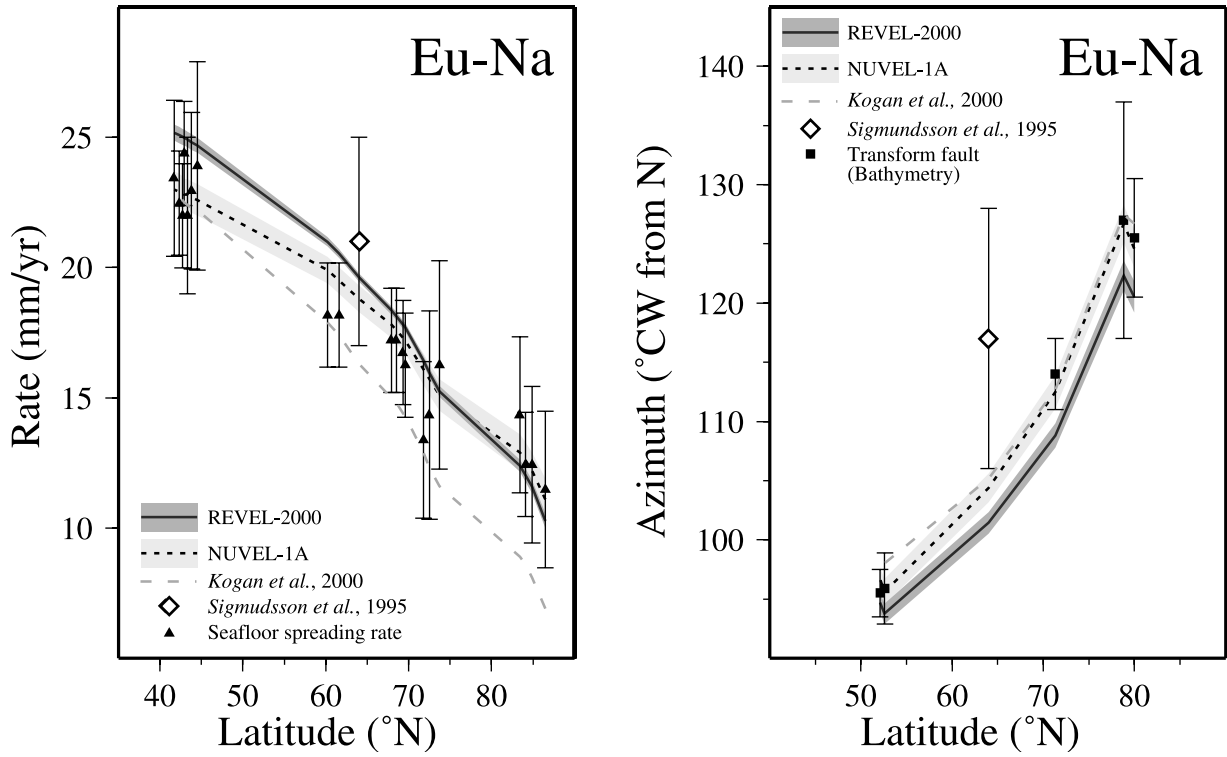
may exist, but our data are insufficient to reliably resolve their relative motion. Given the observed seismicity and surface faulting in the Rhine Graben [Camelbeek and Meghraoui, 1998; Meghraoui *et al.*, 2000], we have taken the more conservative approach and used only sites east of the rift in our definition of Eurasia.

[38] We can nevertheless calculate relative motion across the Rhine Graben, using the admittedly noisy relative angular velocity vector between Asia and western Europe as defined above (58.48°N, 2.35°E, 0.076°/Myr,  $\sigma_{\text{maj}} = 9.4$ ,  $\sigma_{\text{min}} = 2.7$ ,  $\sigma_{\omega} = 0.005$ °/Myr). Calculated at a point on the upper Rhine Graben (49°N, 8°E) we predict west-southwest/east-northeast extension,  $1.5 \pm 0.4$  mm/yr at  $73 \pm 17^\circ$ . Note that this calculation does not require that relative motion between western Europe and Eurasia is accommodated exclusively in the Rhine Graben. Our calculated direction agrees with the orientation of minimum principal stress from earthquake focal mechanisms in the region [Plenefisch and Bonjer, 1997] and is consistent with Nocquet *et al.* [2001]. Westward motion of western Europe may reflect extrusion tectonics associated with the northward motion of Italy and the Nubian plate relative to Eurasia.

[39] Following Kogan *et al.* [2000], we evaluate if site bili in northeastern Russia lies on the North American or Eurasian plate by comparing the size of the corresponding velocity residual. This site has a residual rate of  $6.4 \pm 2.0$  mm/yr with respect to Eurasia and  $3.4 \pm 1.6$  mm/yr with respect to North America, suggesting that bili lies either on the North American plate or on the diffuse boundary between the two plates [Chapman and Solomon, 1976; Zonenshain and Savostin, 1981; Cook *et al.*, 1986].

[40] We can also compare the predictions of our angular velocity estimate for Eurasia-North America with measured spreading rates and transform fault azimuths along the Mid-Atlantic Ridge and with the NUVEL-1A model. Figure 12 shows this comparison and also the predictions of Kogan *et al.* [2000]. REVEL-2000 shows good agreement with the geologic data along the Mid-Atlantic Ridge north of about 65°N but is systematically faster than a cluster of geologic data around 40°–45°N by  $\sim 2$  mm/yr. At 64°N, 20.5°W on the plate boundary in Iceland we predict spreading





**Figure 12.** Similar to Figure 6, for Eurasia-North America.



**Figure 13.** Similar to Figure 5 for the Pacific Ocean area. For clarity, sites on North America are not labeled. Co, Cocos plate.

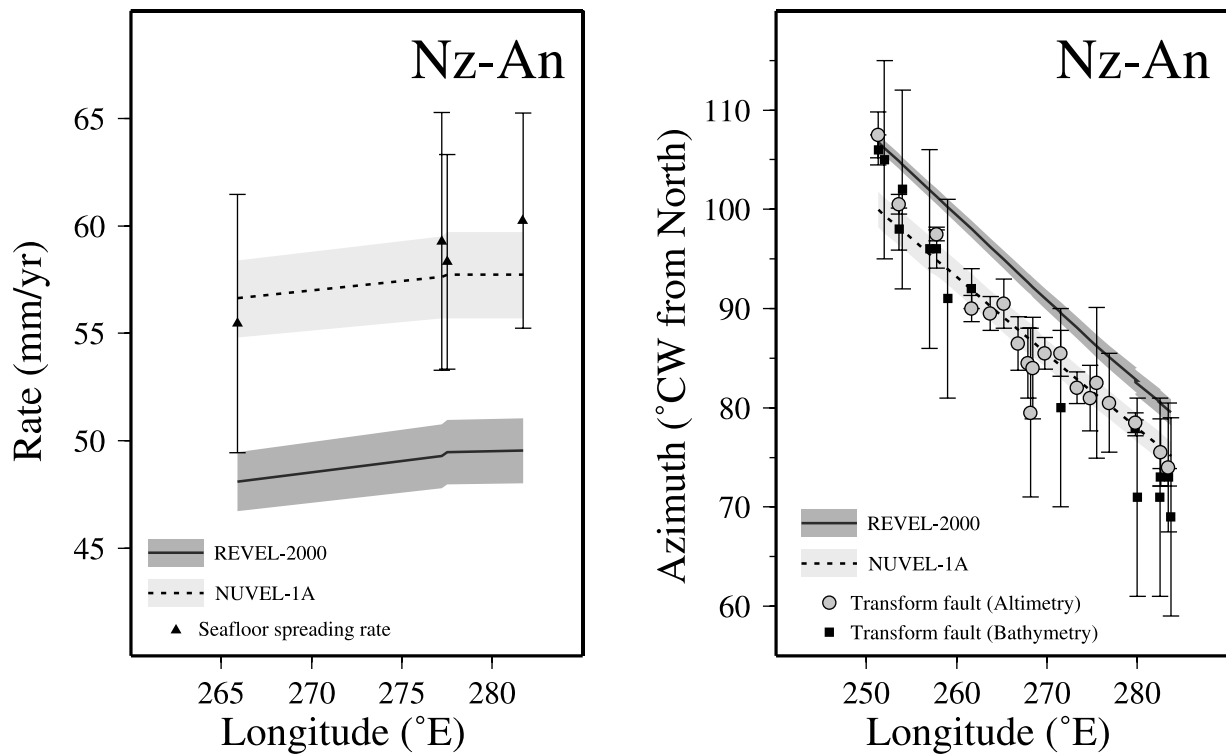


Figure 14. Similar to Figure 6, for Nazca-Antarctica.

at a rate of  $19.9 \pm 0.2$  mm/yr at an azimuth of  $102 \pm 1^\circ$ , which compares well with the local measured velocity using episodic GPS data,  $21 \pm 4$  mm/yr at azimuth of  $117 \pm 11^\circ$  [Sigmundsson *et al.*, 1995].

[41] We can determine what fraction of the total plate motion is accommodated within Iceland by comparing the velocity of hofn in the southeastern corner of the island relative to stable North America ( $21.2 \pm 0.8$  mm/yr at  $104.1 \pm 2.0^\circ$ ) with the relative plate velocity computed at the same location ( $19.7 \pm 0.2$  mm/yr at  $104.2 \pm 0.9^\circ$ ) (hofn is not used in the definition of Eurasia, so this is an independent test). Thus hofn's velocity is consistent with stable Eurasia's velocity within 95% confidence. Similarly reyk in southwestern Iceland is not used in the definition of stable North America and moves at  $20.2 \pm 0.6$  mm/yr at  $281.4 \pm 1.8^\circ$  relative to stable Eurasia, compared to the calculated relative plate velocity here of  $19.9 \pm 0.2$  mm/yr at  $280.8 \pm 0.9^\circ$ , again equivalent within uncertainties. The baseline between hofn and reyk is more precise than the site velocities described above and is independent of any definition of stable North America or Eurasia. Its length rate of change,  $20.3 \pm 0.2$  mm/yr, is very similar to the predicted rate of change from REVEL-2000,  $19.8 \pm 0.2$  mm/yr, calculated at the point where the baseline between these two sites crosses the plate boundary ( $64.2^\circ\text{N}$ ,  $18.8^\circ\text{W}$ ). These results confirm that essentially all of the plate motion between North America and Eurasia is accommodated within the island, consistent with earlier studies [Sigmundsson *et al.*, 1995; Jonsson *et al.*, 1997; Hreinsdottir *et al.*, 2001].

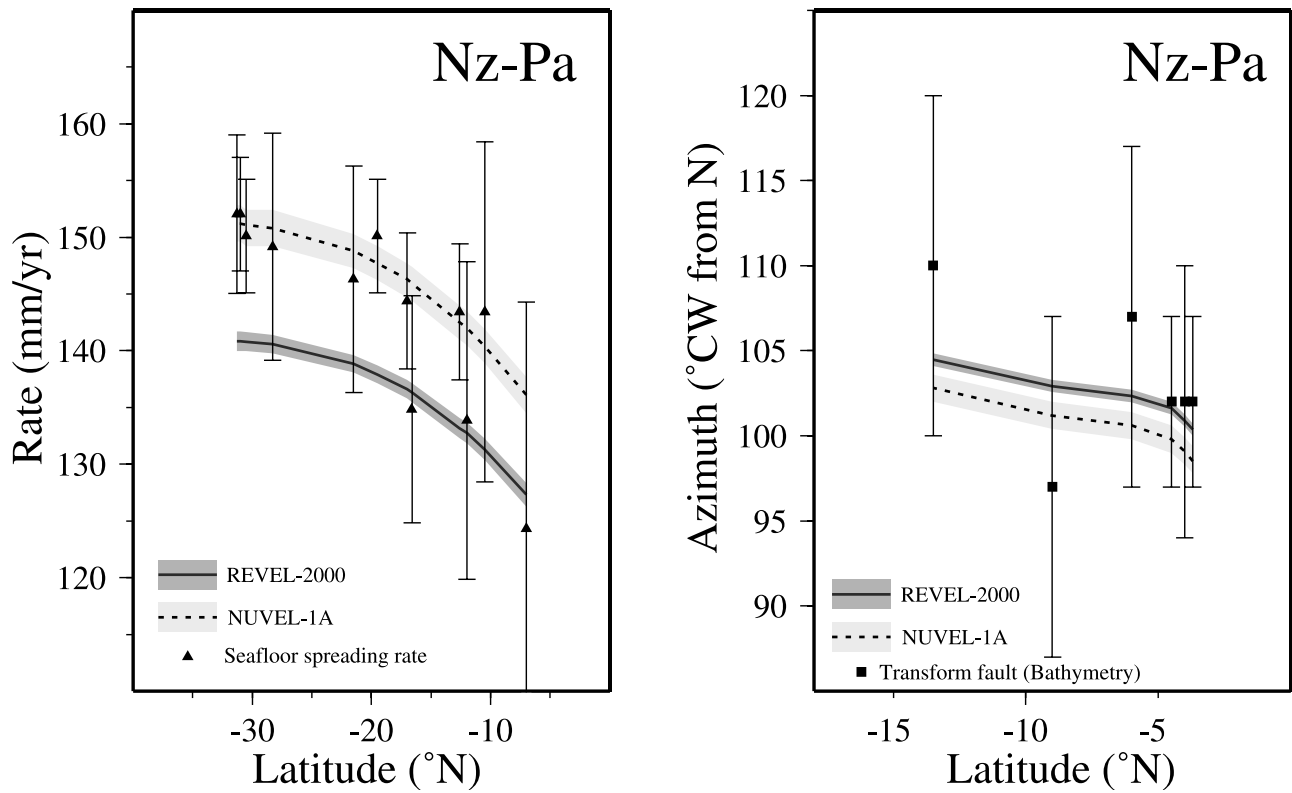
**4.3.7. India.** [42] We define the Indian plate using three sites, HYDE, IISC, and MALD. Our mean rate residual (1 mm/yr) is compatible with earlier rigidity studies of Paul *et al.* [1995] and Malainani *et al.* [2000]. We agree with earlier findings suggesting that the Indian plate is moving slower than predicted by NUVEL-1A [Chen *et al.*, 2000; Shen *et al.*, 2000; Holt *et al.*, 2000; Paul *et al.*, 2001; Kreemer *et al.*, 2000]. These studies show a range of velocity estimates (e.g., rates at IISC relative to Eurasia of 34.8, 41.9, 36, 43.7, and 34.3 mm/yr, respectively), but all are slower than the corresponding NUVEL-1A estimate (47.8 mm/yr). Our Eurasia-India angular velocity predicts a rate of 35.2 mm/yr at this location. It

is unclear whether the difference between the geodetic and geologic rates reflects a true deceleration in the relative velocity of this plate pair or a systematic error in NUVEL-1A [Gordon *et al.*, 1999].

**4.3.8. Nazca and South America.** [43] Owing to the paucity of GPS data on the Nazca plate (EISL, GALA) we include all available space geodetic data, adding SLR (7097) and DORIS (EASA, GALD) (Figure 13). Previous estimates for South America were also hampered by relatively sparse data. For example, Norabuena *et al.* [1998] had only four continuous GPS sites with a total of  $\sim 3000$  station days of data to define stable South America, while Norabuena *et al.* [1999] had 6200 station days of data, also at four sites. For this study we have 11 GPS sites with 14,200 station days of data; hence South America's velocity is much better defined. Our resulting Nazca-South America rate is nevertheless consistent with previous results [Norabuena *et al.*, 1998, 1999; Angermann *et al.*, 1999] and significantly slower than the NUVEL-1A estimate. Nazca-Antarctica and Nazca-Pacific are also significantly slower than the geologic model (Figures 14 and 15). As has been shown in previous studies, both Nazca-Pacific and Nazca-South America have been decelerating during the last 25 Myr [Tebbens and Cande, 1997; Somoza, 1998]. A similar deceleration applies to Nazca-Antarctica (Figure 16). These decelerations are sufficiently rapid that they can be observed as differences between geodetic plate motion estimates and NUVEL-1A predictions.

[44] The REVEL-2000 velocity azimuths for Nazca-Antarctica and Nazca-Pacific also differ from NUVEL-1A (Figure 16). Although the Nazca site distribution is limited, the difference may be real, perhaps reflecting changing plate direction over time. Stage pole reconstructions back to 30 Ma [Tebbens and Cande, 1997] allow us to look at a much longer time record and suggest changes in Nazca-Antarctica and Nazca-Pacific direction in the last 15 Myr in the same sense as we infer from the shorter epochs "sampled" by NUVEL-1A and REVEL-2000 (compare Figures 14 and 16).

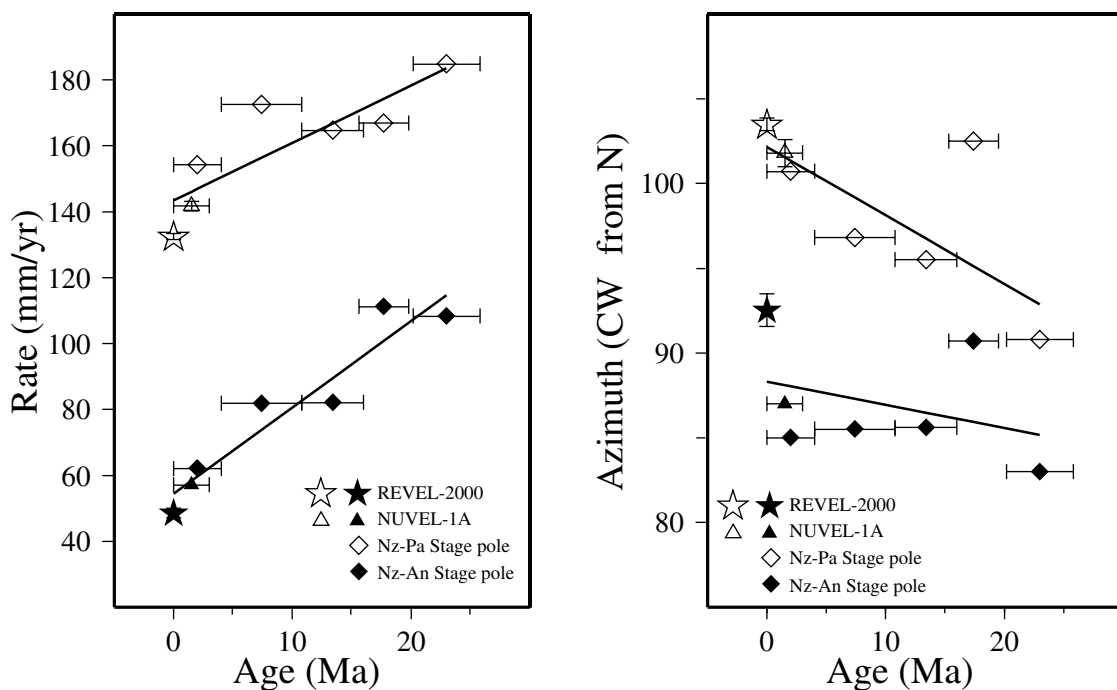
**4.3.9. North America.** [45] The rigid North American plate is defined using 64 sites (Figure 13 and Table 1). We exclude all sites located in the Gulf Coast because of possible subsidence and



**Figure 15.** Similar to Figure 6, for Nazca-Pacific.

all sites west of the Rio Grande Rift and west of the Rocky Mountain front because of possible tectonic effects. Sites mem2 in Memphis, Tennessee, and cha1 in Charleston, South Carolina, are also omitted as they have both been the focus of major intraplate

earthquakes in the last 200 years [Nutli, 1973; Bollinger, 1972]. As discussed earlier, we omit 20 sites likely to be most affected by glacial isostatic adjustment in the northern part of the plate. A more detailed discussion of North America is in preparation.



**Figure 16.** Spreading rates and transform azimuths predicted for Nazca-Pacific (calculated at 12°S, 250°E) and Nazca-Antarctic (calculated at 40°S, 268°E) plates using NUVEL-1A, REVEL-2000, and stage pole data of *Tebbens and Cande* [1997]. Note more rapid deceleration implied by post-10 Ma data.

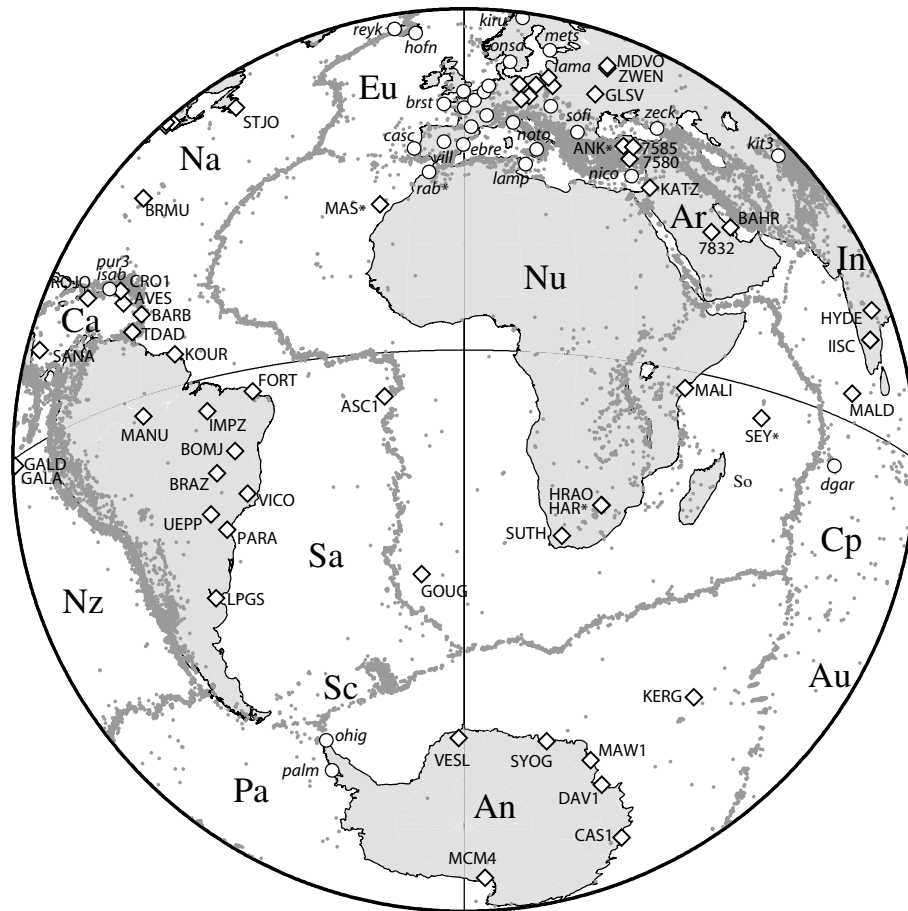


Figure 17. Similar to Figure 5 for the Atlantic Ocean area. Sc, Scotia plate.

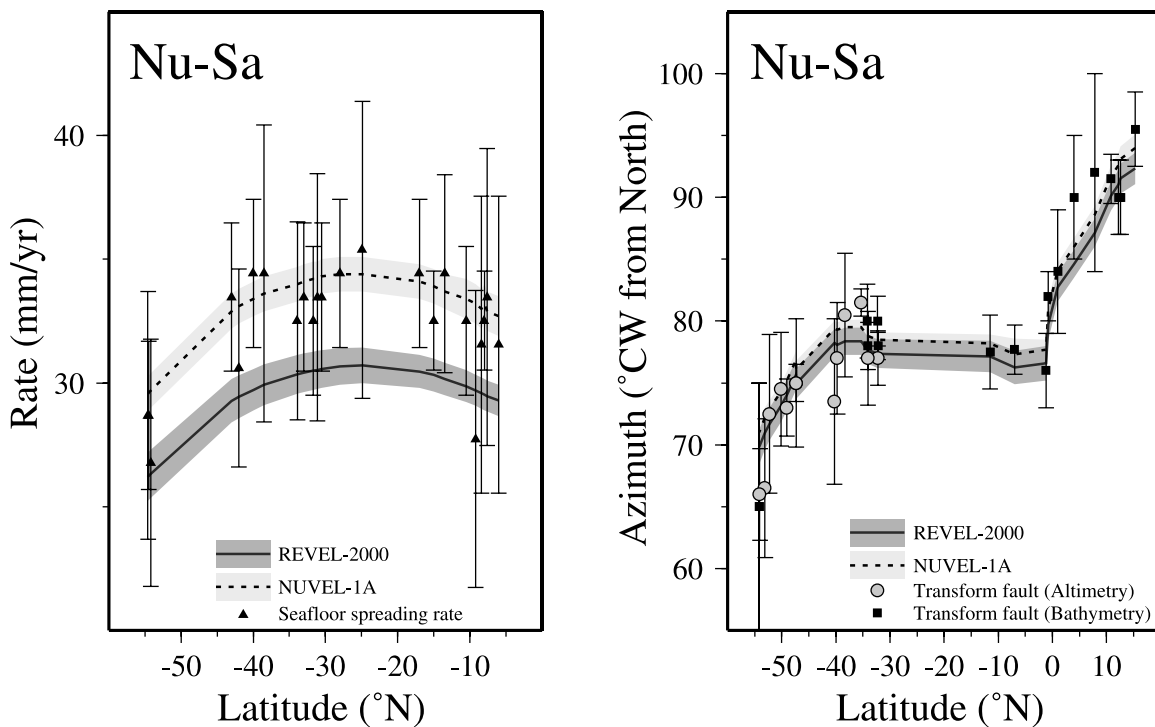
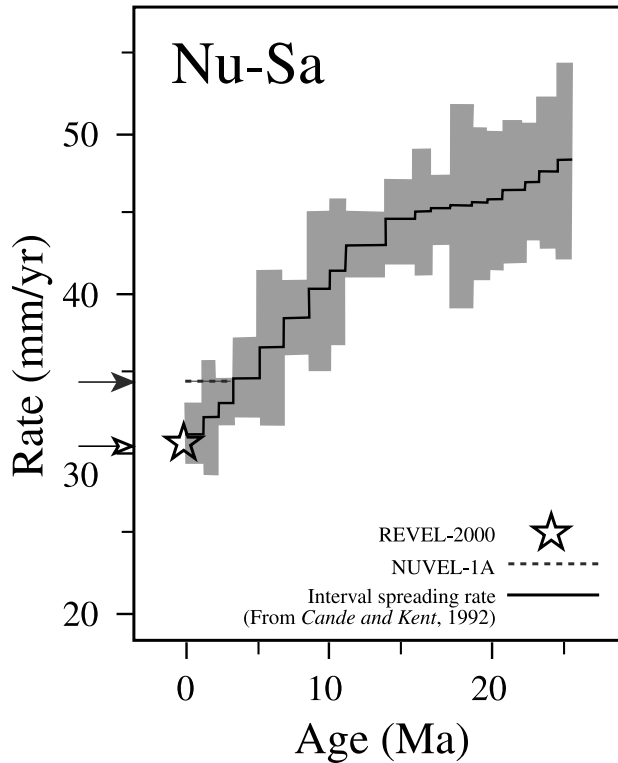


Figure 18. Similar to Figure 6, for Nubia-South America.





**Figure 19.** Nubia-South America spreading rates (full rate) versus age for last 25 Myr, from *Cande and Kent* [1992] (shaded area shows  $\sim 1$  standard error), compared to NUVEL-1A and REVEL-2000 predictions (arrows) calculated at  $28^{\circ}\text{S}$ ,  $348^{\circ}\text{E}$ . REVEL-2000 differs significantly from NUVEL-1A but agrees closely with an extrapolation of interval spreading rates to Recent time.

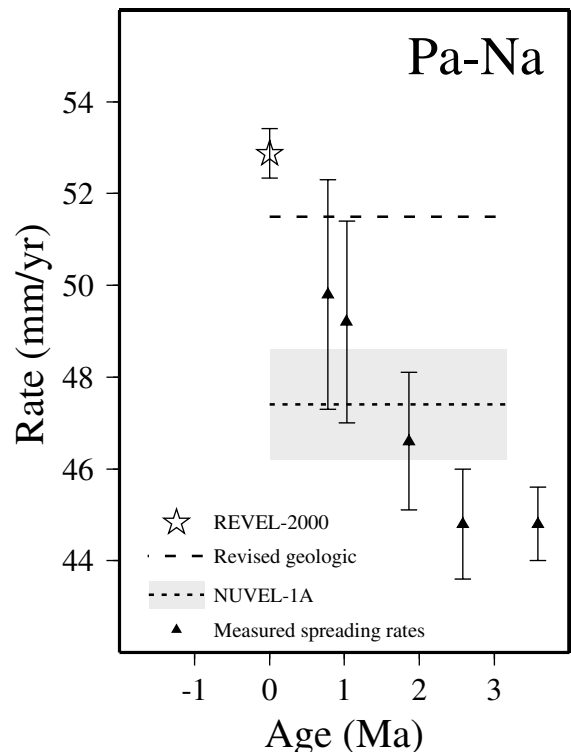
[46] Fairbanks, Alaska (fair), has been used in our previous definitions of stable North America [e.g., *Dixon et al.*, 1996]. It is excluded here because it is too close to the seismically active plate boundary zone and moves  $2.2 \pm 0.5$  mm/yr at  $171 \pm 11.4^{\circ}$  relative to stable North America (Table 1), somewhat faster than would be expected for a “stable plate” interior site.

**4.3.10. Nubia.** [47] We define an angular velocity for Nubia using five sites (Goug, HAR\*, HRAO, MAS\*, SUTH). HRAO and HAR\* in South Africa are located near an area of intense microseismicity (Figure 17). A three-site solution that omits these two sites has an angular velocity of  $52.65^{\circ}\text{N}$ ,  $-82.79^{\circ}\text{E}$ ,  $0.249^{\circ}/\text{Myr}$ ,  $\sigma_{\text{max}} = 1.9$ ,  $\sigma_{\text{min}} = 0.6$ ,  $\sigma_{\omega} = 0.004$ ,  $\chi^2_{\nu} = 0.27$ ,  $\text{MRR} = 0.79$  mm/yr, very similar to our preferred five-site solution (Table 2), and suggests that sites HRAO and HAR\* can be included in the stable plate solution.

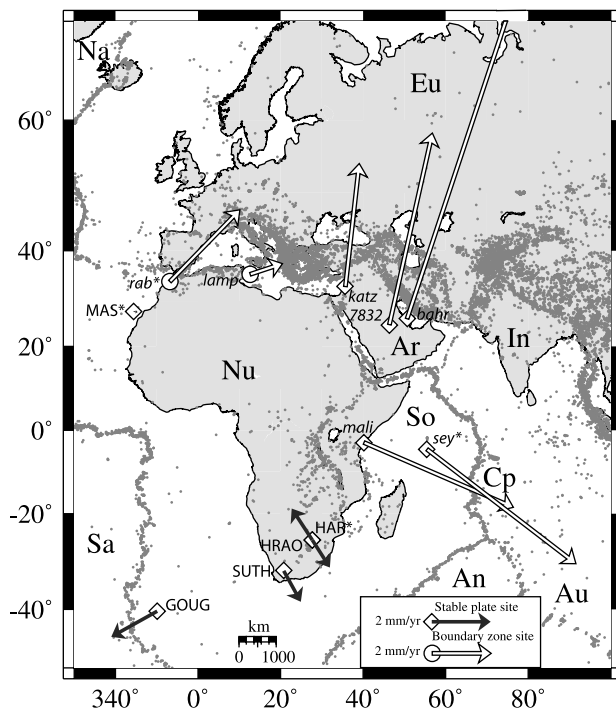
[48] Our predicted Nubia-South America velocity should not be directly comparable to the NUVEL-1A model, since NUVEL-1A is based on a composite African plate, while we define separate Nubia and Somalia plates. However, we can still compare with the underlying geologic data (Table 5). Despite this difference, REVEL-2000 and NUVEL-1A azimuths in the South Atlantic are indistinguishable from each other and from the geologic data (Figure 18). However, the REVEL-2000 and NUVEL-1A rates differ significantly, with REVEL-2000 slower than both NUVEL-1A and the geologic data upon which it is based, by  $\sim 4$  mm/yr through a large range of latitudes (Figure 18). The simplest interpretation is that this reflects true deceleration rather than a systematic bias in either model, since both models are well constrained in this region. Longer-term geologic data support this inference. Figure 19, modified from *Cande and Kent* [1992], shows a remarkable

agreement between the REVEL-2000 geodetic rate and a longer-term trend of decelerating spreading in the south Atlantic going back  $\sim 25$  Myr. This is roughly the time of initiation of the current phase of Andean crustal shortening (see discussion of slowing Nazca-South America convergence by *Norabuena et al.* [1999, and references therein]). It is tempting to speculate that the same process that has slowed convergence between South America and Nazca (e.g., formation of the Andes, associated crustal thickening, and possible increased resistance to subduction) may also contribute to a gradual slowing of South America’s westward component of motion and consequent slowing of spreading in the south Atlantic. For example, growth of the Andes and high topography would increase South America’s east directed gravitational body force, opposing the west directed body force associated with spreading at the Mid-Atlantic Ridge.

**4.3.11. Pacific.** [49] We define the Pacific plate using 9 sites (Table 1 and Figure 13), excluding two sites (mkea and upo1) on the Big Island of Hawaii that may be affected by deformation associated with the active Mauna Loa and Kilauea volcanoes. We agree with the conclusions of *DeMets and Dixon* [1999] that the NUVEL-1A angular velocity for Pacific-North America significantly underestimates present-day Pacific-North American motion. *DeMets* [1995] discussed possible biases in the NUVEL-1A estimate for Pacific-North America, in particular, the rate data based on magnetic anomalies from the Alarcon Rise in the southern Gulf of California, which do not reflect the full plate motion due to faults to the west that have been active for most of the last 3 Myr. These may also be active today [*Dixon et al.*, 2000b]. Thus the bias in NUVEL-1A’s estimate of Pacific-North America motion has essentially the same explanation as biases in Caribbean-North America and Caribbean-South America motion, namely, incorporation of rate data in the geologic model that, because of tectonic complexity, do not reflect total plate motion.



**Figure 20.** Measured seafloor spreading rate at various times in the southern Gulf of California corrected for outward displacement, from *DeMets* [1995], and predicted total motion from REVEL-2000 calculated at  $23.5^{\circ}\text{N}$ ,  $251.5^{\circ}\text{E}$ . Revised geologic model from *DeMets and Dixon* [1999].



**Figure 21.** Similar to Figure 10, for the Nubian plate. Also shown are the velocity vectors for two sites on the Somalia plate (mali, sey\*) and three sites on the Arabian plate (bahr, katz, 7832) relative to stable Nubia, indicating extension across the East African Rift, Red Sea, and Gulf of Aden. For clarity, error ellipses are omitted.

[50] Spreading in the southern Gulf of California shows a trend of increasing rate since 2.5 Ma [DeMets, 1995], reflecting a combination of increased “focusing” of plate motion as spreading at the Alarcon Rise increasingly reflects the full plate motion and perhaps a small amount of acceleration. REVEL-2000’s estimate of Pacific-North America motion is  $1.4 \pm 0.5$  mm/yr faster than the geologic estimate of DeMets and Dixon [1999] for the southern Gulf of California (Figure 20).

[51] Six Pacific sites near the plate boundary zone with North America (cicz, farb, scip, sni1, spmx, vndp) have velocities slower than expected if they were on the rigid Pacific plate. Dixon *et al.* [2000b] and Beavan *et al.* [2002] explain these discrepancies by a combination of strain accumulation on locked faults of the San Andreas system, plus slip on additional faults offshore to the west [e.g., Sorlien *et al.*, 1999]. Figure 20 suggests that such faults in the region of southern Baja California have a total slip rate less than  $\sim 2$  mm/yr (the difference between the REVEL-2000 rate and estimated present-day spreading rate in the Alarcon Rise based on a linear fit to the magnetic anomaly rate data younger than 3 Ma).

[52] We are unable to directly verify our Australia-Pacific angular velocity against independent geologic data owing to the lack of transform faults or spreading ridges. Australia-Pacific stage pole data suggest little or no change over the last 11 Myr [Sutherland, 1995]. At a point on the Alpine fault in New Zealand ( $43.5^\circ\text{S}$ ,  $170^\circ\text{E}$ ), REVEL-2000 predicts  $42.0 \pm 0.6$  mm/yr at an azimuth of  $248 \pm 1^\circ$ , giving  $41 \pm 1$  mm/yr of fault-parallel slip (azimuth  $236^\circ$ ) and  $8 \pm 1$  mm/yr of shortening. The fault-parallel component is larger than the rate measured with episodic GPS, although the fault-normal component is comparable: Beavan *et al.* [1999] obtain  $29.7 \pm 1.4$  mm/yr of fault-parallel slip and  $9.9 \pm 1.8$  mm/yr of shortening. Beavan *et al.* [1999] and Sutherland *et al.*

[2000] conclude that the Alpine fault and subparallel faults in New Zealand take up only  $\sim 70$ – $80\%$  of total plate motion.

**4.3.12. Philippine.** [53] Angular velocity estimates for the Philippine plate by conventional geologic approaches are problematic because no spreading ridges bound the plate, and therefore geologic rate data are not available. This problem has been attacked with earthquake slip vector data [Seno *et al.*, 1993] and GPS data [Kato *et al.*, 1996, 1998; Kotake *et al.*, 1998]. We use four sites (GSI1, GSI3, OKTO, and PAL\*);  $\chi_v^2 = 1.03$  to define the rigid Philippine plate (Figure 5). To test if sites close to the subducting Pacific plate experience strain accumulation or other nonrigid plate effects, we calculated residual velocities for ccjm, gsi2, and haci with respect to the rigid Philippine plate as defined above. Sites ccjm and gsi2, both  $\sim 100$  km from the trench, have residual velocities of  $5.0 \pm 1.7$  mm/yr at  $274 \pm 16^\circ$  and  $3.0 \pm 1.8$  mm/yr at  $142 \pm 32^\circ$ , respectively. Site haci, 200 km from the trench, but within 100 km of Amuria (mainland Japan), has a residual of  $11.1 \pm 1.7$  mm/yr at  $119.9 \pm 5.4^\circ$ . The eastward direction of the residuals at gsi2 and haci suggests that these sites are not strongly influenced by subduction strain accumulation from a seismically coupled Pacific plate (otherwise they would move west). Their eastward velocities may represent postseismic effects, or more likely slow spreading of a back arc basin behind the subduction zone. We cannot exclude the possibility that both effects occur, are of opposite sign, and sum to give the observed residual velocity. Site ccjm’s westward residual velocity may reflect seismic coupling. Inclusion of any of these sites increases misfit ( $4 + \text{ccjm}$ ,  $\chi_v^2 = 2.05$ ;  $4 + \text{gsi2}$ ,  $\chi_v^2 = 1.19$ ;  $4 + \text{haci}$ ,  $\chi_v^2 = 7.01$ ). Site s102, located just offshore Taiwan, clearly does not represent the stable interior of the Philippine plate ( $4 + \text{s102}$ ,  $\chi_v^2 = 21.5$ ) in agreement with Yu *et al.* [1997, 1999].

[54] The velocity of guam with respect to the Philippine plate is a measure of spreading across the Mariana Trough, a back arc basin known to be actively extending. We predict a spreading velocity across the Mariana Trough at guam ( $13.6^\circ\text{N}$ ,  $144.9^\circ\text{E}$ ) of  $46.6 \pm 1.5$  mm/yr at an azimuth of  $096.3 \pm 1.8^\circ$ , in broad agreement with published geologic rates (full spreading rate), 30–60 mm/yr [Bibee *et al.*, 1980; Hussong and Uyeda, 1981; Ishihara *et al.*, 2001]. Part of this range may reflect a change in spreading rate with latitude, with rate increasing to the south [Stern *et al.*, 1984]. The subduction velocity of the Pacific plate relative to the overlying plate at the southern Mariana trench is defined by guam’s velocity relative to the Pacific,  $63.4 \pm 1.1$  mm/yr at an azimuth of  $104.9 \pm 0.7^\circ$ . The velocity of guam could be affected by strain accumulation from the subduction zone; thus the subduction rate estimate is a lower limit. However, seismic coupling in this region is believed to be low owing to the old age of the subducting plate [Uyeda and Kanamori, 1979].

[55] A similar analysis, and with the same caveats, can be applied to the western margin of the Philippine plate along the central Ryukyu arc, comparing Philippine and South China angular velocities with site oknw. The subduction rate of the Philippine plate in the central Okinawa Trough is defined by oknw’s velocity relative to the Philippine plate,  $98.1 \pm 1.3$  mm/yr at  $131.9 \pm 0.8^\circ$ . We predict a rate of spreading across the back arc basin at oknw of  $21.4 \pm 1.4$  mm/yr at  $173.7 \pm 4.1^\circ$  using our South China angular velocity. This is consistent with estimates of extension from seismic reflection data, 10–20 mm/yr [Park *et al.*, 1998], and earthquake slip vectors azimuths ( $\sim 150^\circ$ ) in the central part of the Okinawa Trough [Fabbri and Fournier, 1999].

**4.3.13. Somalia.** [56] For the Somalia (East Africa) plate we have only two sites (MALI and SEY\*) (Figure 17). SEY\* is site SEY1 omitting the first year of data that is particularly noisy (Table 1) (this is the only example in this study where data were excluded). The resulting Somalia-Nubia angular velocity predicts extension across the East African Rift in an east-southeast to west-northwest direction, at rates and azimuths that are in

approximate agreement with geologically based estimates [Jestin *et al.*, 1994; Chu and Gordon, 1999] (Figure 21 and Table 5). For example, at the equator on the active western rift (30°E) we predict  $5.5 \pm 0.4$  mm/yr of extension at an azimuth of  $98.3 \pm 4.3^\circ$ . At 10°N, 40°E, we predict  $6.9 \pm 0.4$  mm/yr of extension at  $107.9 \pm 3.7^\circ$ .

**4.3.14. South China.** [57] Three sites are used to define the angular velocity of the South China block: S012, SHAO, and WUHN (Figure 5). We can estimate the relative motion of adjoining plates or blocks by determining the velocity of selected sites with respect to South China or by predicting relative velocities across faults based on our angular velocity, for comparison to geologic data. Site kunn located southwest of the South China block, across the north-south striking, left-lateral Xiaojiang fault forming the block's southwest boundary, has a velocity of  $10.2 \pm 2.2$  mm/yr with an azimuth of  $200 \pm 16^\circ$ . This is consistent with the geologic estimate of  $5 \pm 3$  mm/yr for the Xiaojiang fault [England and Molnar, 1997] and with GPS measurements of Chen *et al.* [2000] that suggest  $\sim 10$  mm/yr motion across the fault. The east-west striking left-lateral Qinling fault defines the northern boundary with Amurian plate. At a point on the western part of the fault (34.5°N, 109°E) our Amuria-South China angular velocity predicts  $4.1 \pm 1.4$  mm/yr at  $260 \pm 49^\circ$ , compatible with geologic observations [Zhang *et al.*, 1995] showing  $7.2 \pm 2.2$  mm/yr of motion. On the eastern part of the fault (30°N, 116°E) we predict  $3.4 \pm 2.5$  mm/yr at  $278 \pm 33^\circ$ , where Zhang *et al.* [1995] measure  $2 \pm 1$  mm/yr.

## 5. Conclusions

[58] We have derived a kinematic model for the motion of 19 lithospheric plates and continental blocks based on high-precision space geodetic data. By excluding sites that may be influenced by seismic cycle effects within the plate boundary zone as well as sites affected by glacial isostatic adjustment, we believe the plate velocity model to be representative of geologically Recent motions (last  $\sim 10,000$  years) and have termed it REVEL, for Recent velocities. Our model includes a rigorous and independent estimate for GPS velocity errors, facilitating tests of plate rigidity and comparison with plate velocities for different time periods. Departures from short-term rigid plate behavior due to glacial isostatic adjustment are clearly observed for North America, Eurasia, and Antarctica. Possible departures from plate rigidity due to longer-term intraplate deformation are observed in Australia and Eurasia; the latter can be interpreted to indicate extension across the Rhine Graben. We see statistically significant differences between the velocity predictions of REVEL-2000 and those of the NUVEL-1A geologic model for about one third of tested plate pairs. Pacific-North America motion and motion of the Caribbean plate with respect to North and South America are significantly faster than NUVEL-1A, presumably reflecting systematic errors in the geologic model because the relevant rate data do not reflect the full plate rate. Many other differences appear to reflect real velocity changes over the last few million years. Arabia-Nubia and Arabia-Eurasia may be slowing, perhaps in response to the collision of Arabia with Eurasia and consequent increased resistance to Arabia's northward motion. Nazca-Pacific, Nazca-Antarctica, Nazca-South America, and South America-Nubia are slower than the 3-Myr average, part of a longer-term slowing trend dating back 25–30 Myr, possibly associated with ongoing construction of the Andes [Norabuena *et al.*, 1999].

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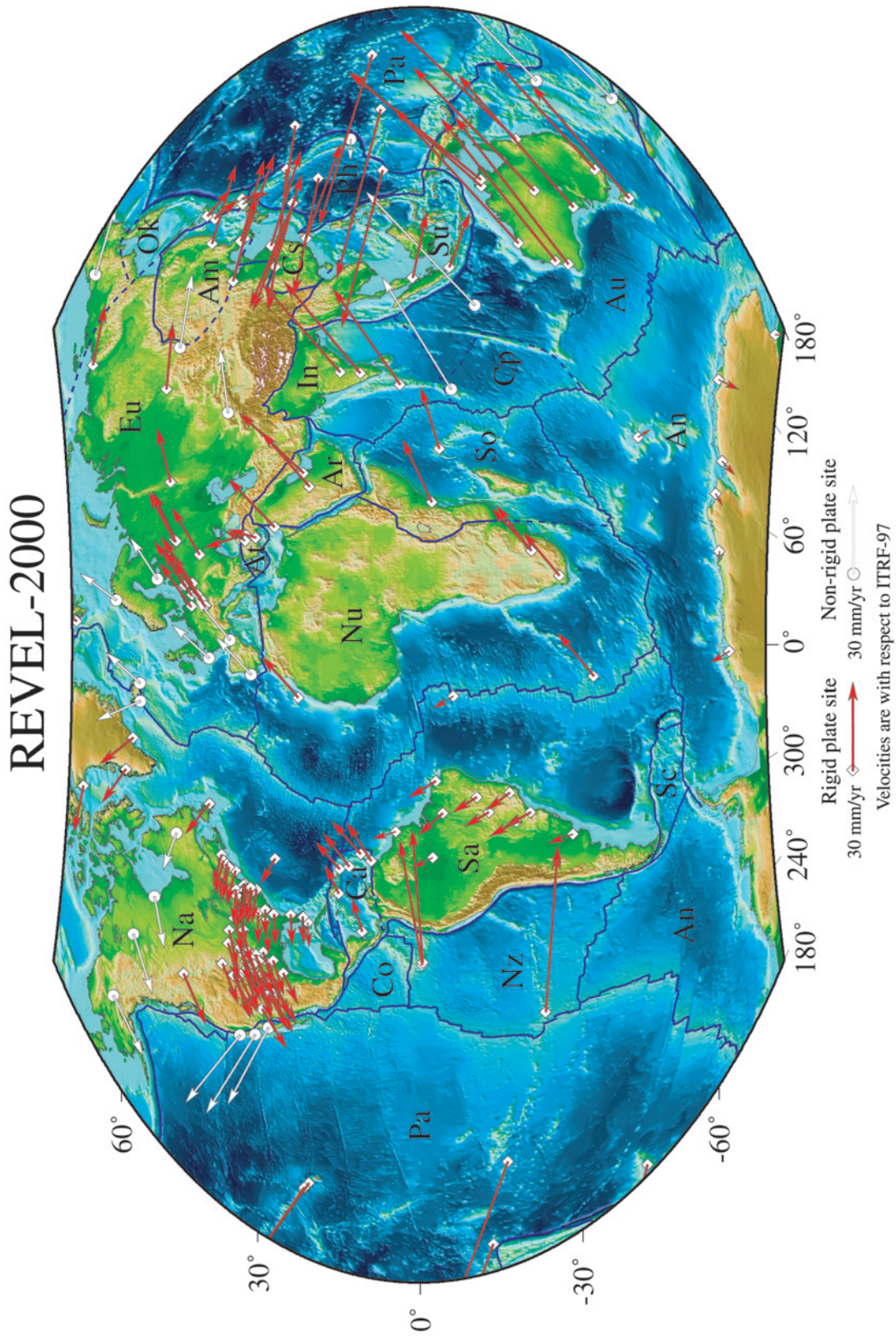
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**Figure 4.** Winkel Tripel projection showing site velocities with respect to ITRF-97. Solid arrows (red) with diamonds are sites on rigid plate interiors; for clarity, sites on the Sierra Nevada block are omitted. Open arrows (white) with circles are selected sites in deforming plate boundary zones, plate interiors affected by GIA or other locations such that they are not used in the rigid plate definition. Plate boundaries are shown with blue lines, dashed where plate boundary is poorly defined. See Table 1 for plate abbreviations, except Sc, Scotia, and Cp, Capricorn.